Determining atmospheric aerosol content with an infra-red radiometer

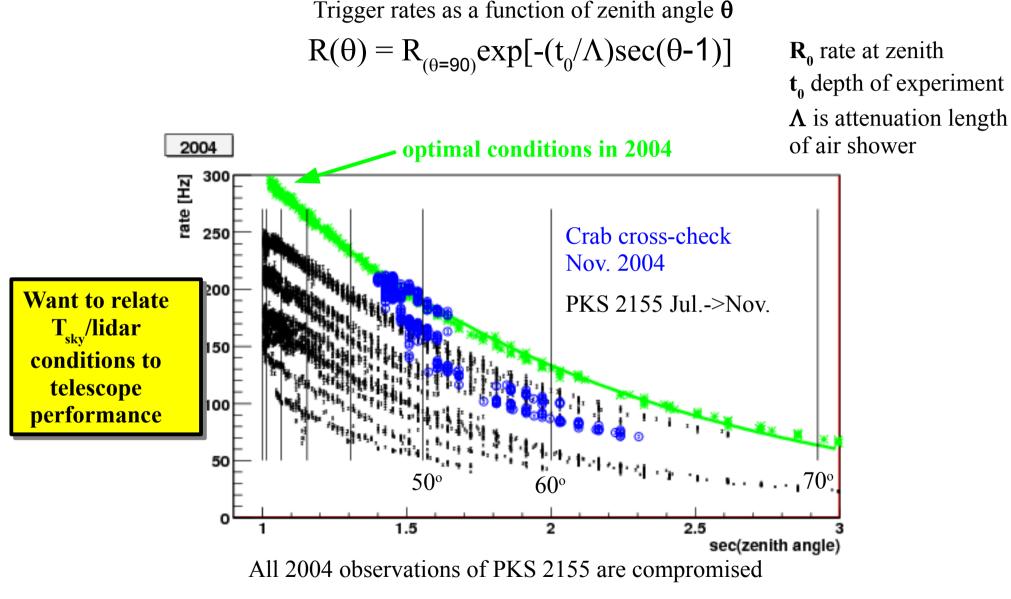
Michael Daniel (michael.daniel@liverpool.ac.uk) Y.T.E. Lo P.M. Chadwick



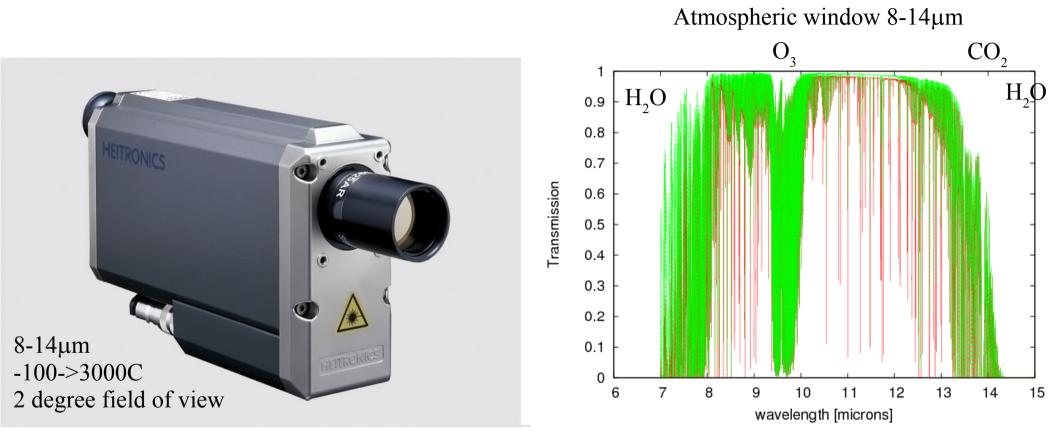




Motivation: the recovery of compromised observing runs, eg taken as part of a multiwavelength campaign



Heitronics KT19.82 IR radiometer



Gemini North, Lock (1992).

Remember, the radiometer is not measuring *temperature*, it is measuring radiance

$$L_{sky} = \sigma T_{sky}^{4}$$

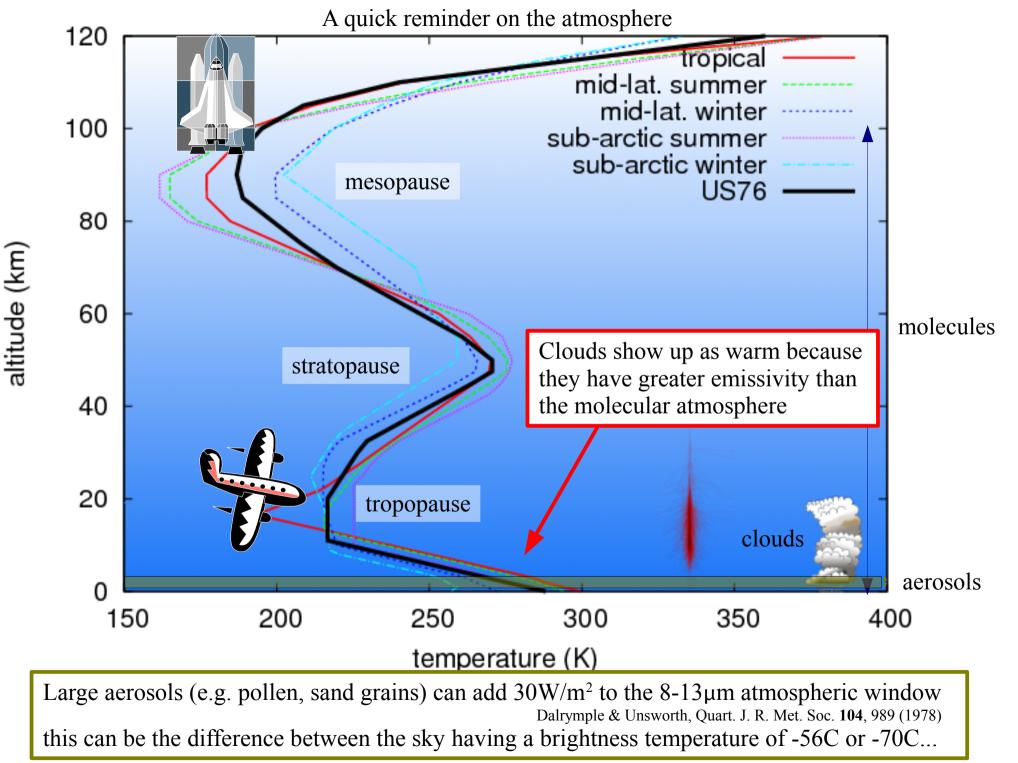
= $L_{window} + L_{H_2O} + L_{aerosol} + L_{molecular} + ...$
= $\epsilon_w \sigma T_{window}^{4} + \epsilon_{wv} \sigma T_{H_2O}^{4} + \epsilon_a \sigma T_{aerosol}^{4} + \epsilon_m \sigma T_{molecular}^{4} + ...$

this is not a linear function of temperature

The efficient emitters and the warmest components are going to dominate the effective sky brightness temperature

We just need to work out how best to make use of the sensitivity

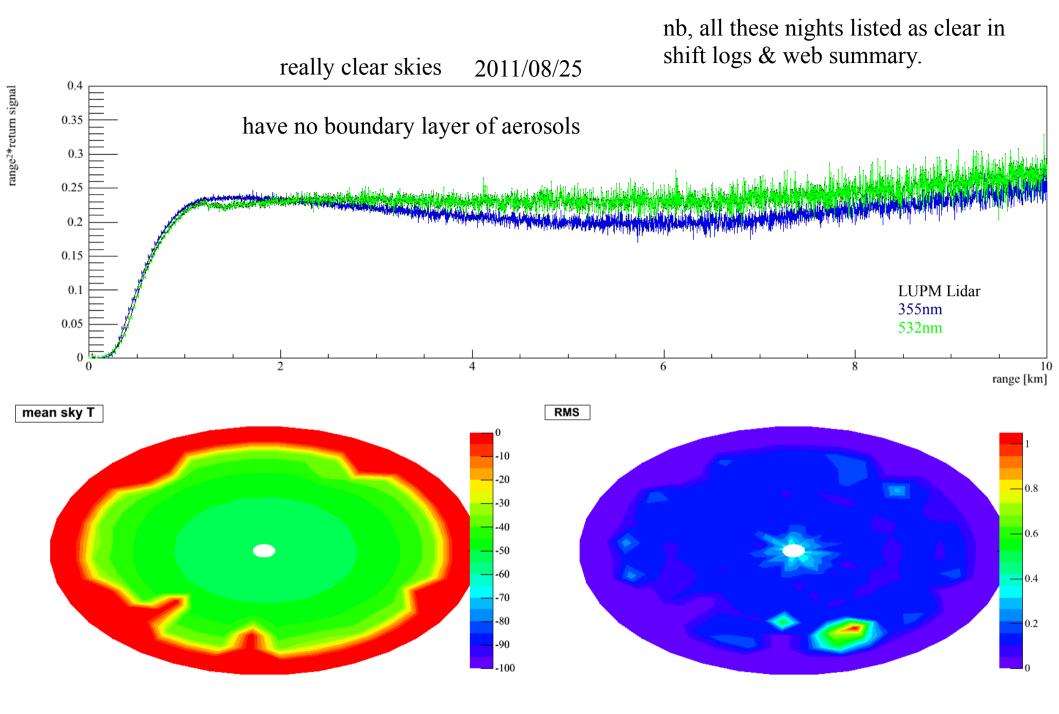
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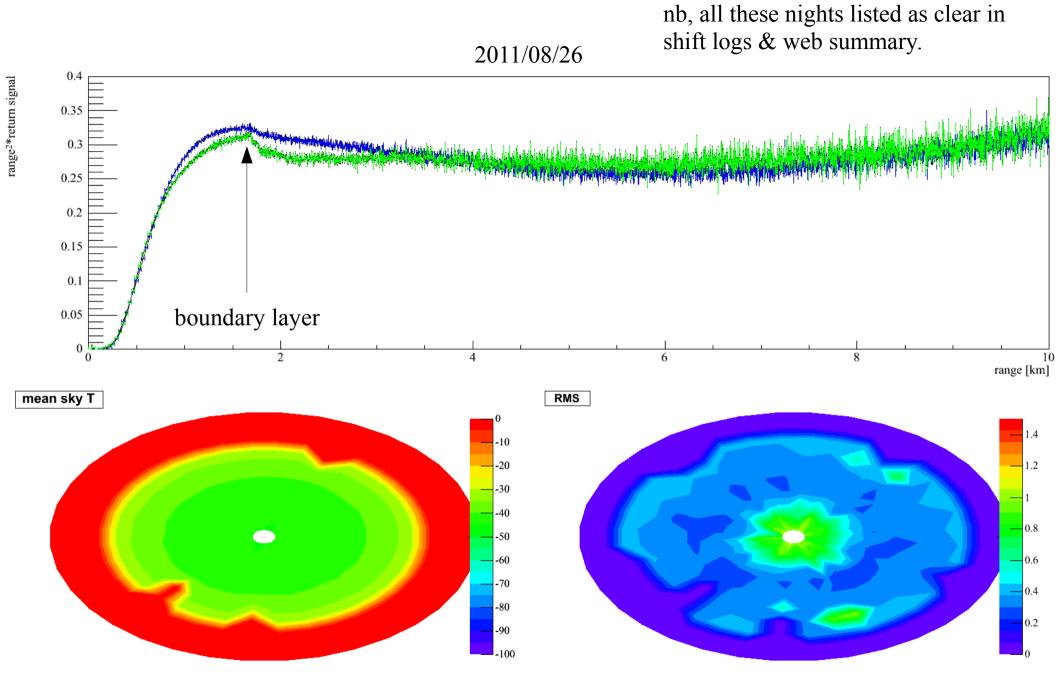


Effect of aerosol contribution on sky brightness temperature

Large aerosols (e.g. pollen, sand grains) can add 30W/m² to the 8-13µm atmospheric window Dalrymple & Unsworth, Quart. J. R. Met. Soc. **104**, 989 (1978)

molecular luminosity [W/m²]	brightness temperature [C]	T with additional 30 W/m ² [C]
140	-50	-39
117	-60	-47
97	-70	-56
79	-80	-64

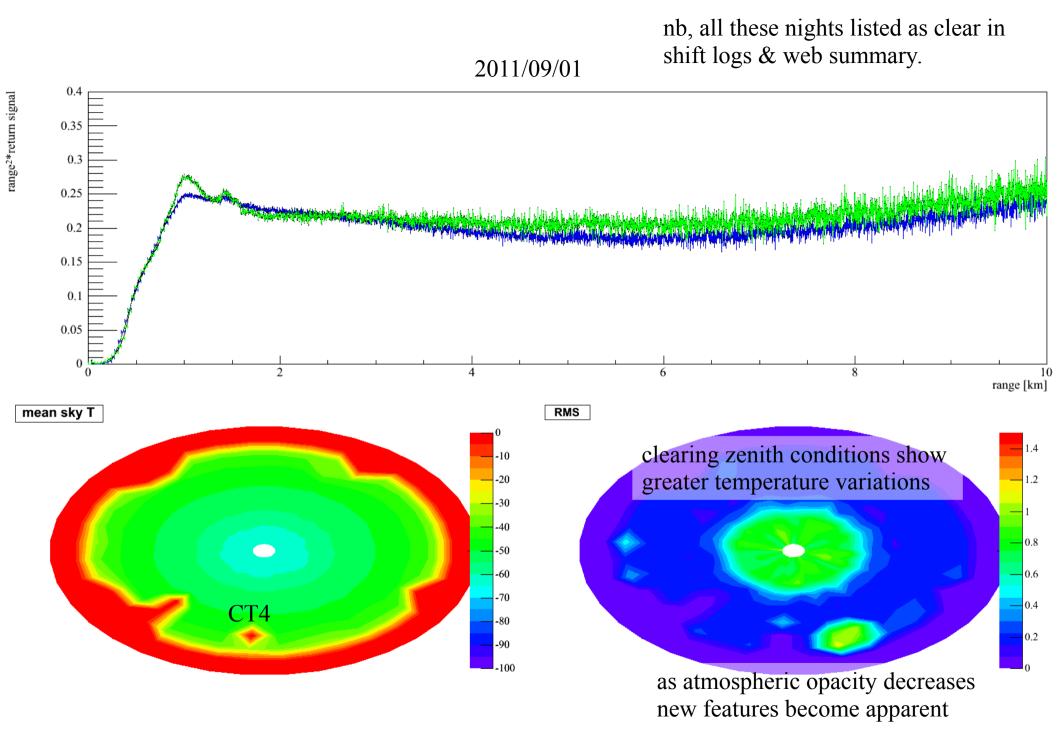




shift logs & web summary. "normal" observing conditions 2011/08/29 0.4 0.35 0.3 0.25 obvious boundary layer 0.2 0.15 0.1 0.05 0 L 0 2 10 atmosphere is also more opaque in the IR range [km] mean sky T RMS 1.4 -10 -20 1.2 -30 -40 0.8 -50 -60 0.6 -70 0.4 -80 0.2 -90 -100

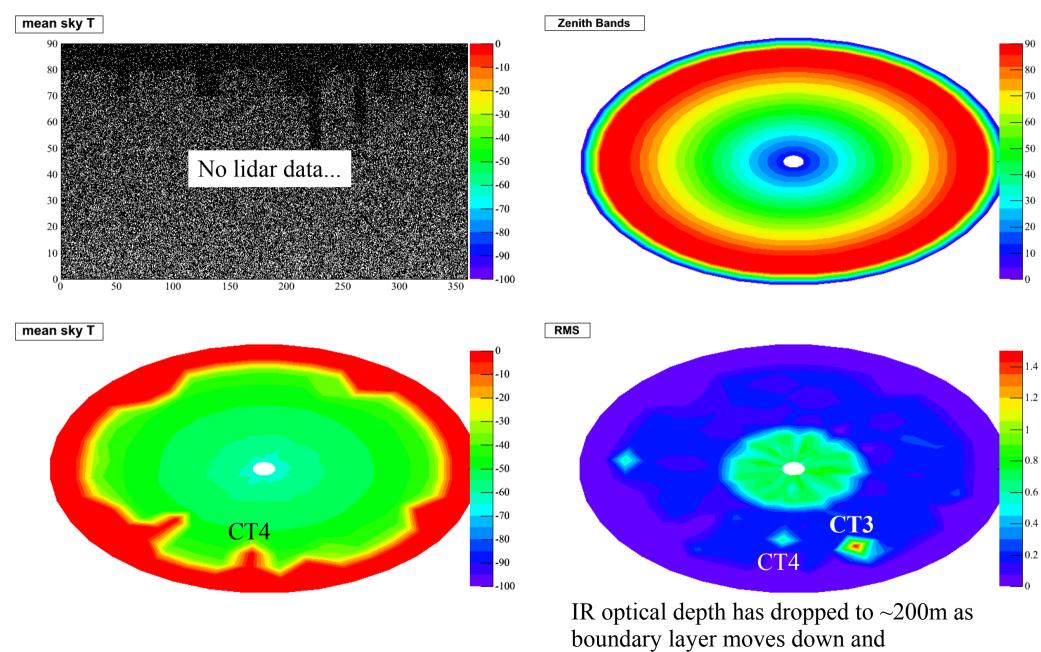
range2*return signal

nb, all these nights listed as clear in



nb, all these nights listed as clear in shift logs & web summary.

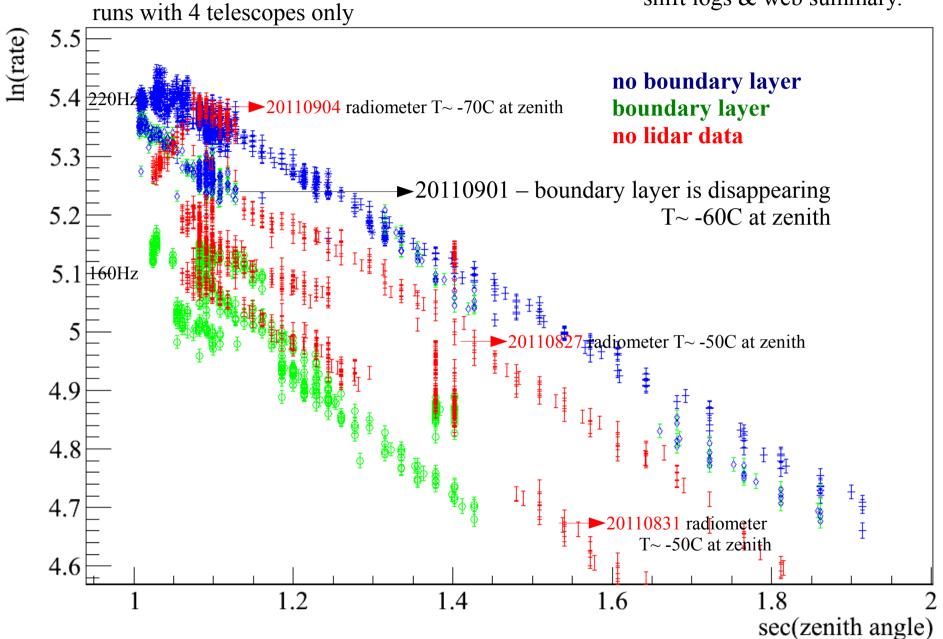
2011/09/04



aerosols have sedimented out of atmosphere

comparing those nights telescope trigger rates

nb, all these nights listed as clear in shift logs & web summary.



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Extending the study

H.E.S.S. July 2011 – July 2012

To provide a comparison with telescope trigger rates & lidar observations

Clear sky:

– no obvious clouds, haze or aerosols

- at least 7 runs with $>35^{\circ}$ spread in zenith angles
- full 28 minute observing run (no hardware, or sun/moon rise contamination)
- lidar run before/after shows no developed aerosol boundary layer

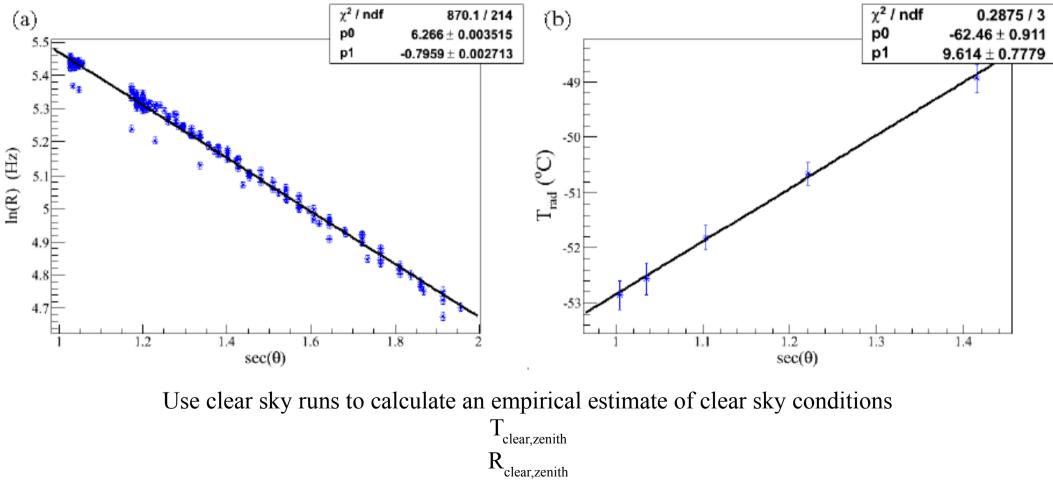
Leads to 19 'good' nights (~100 runs with lidar data) concentrated July->December

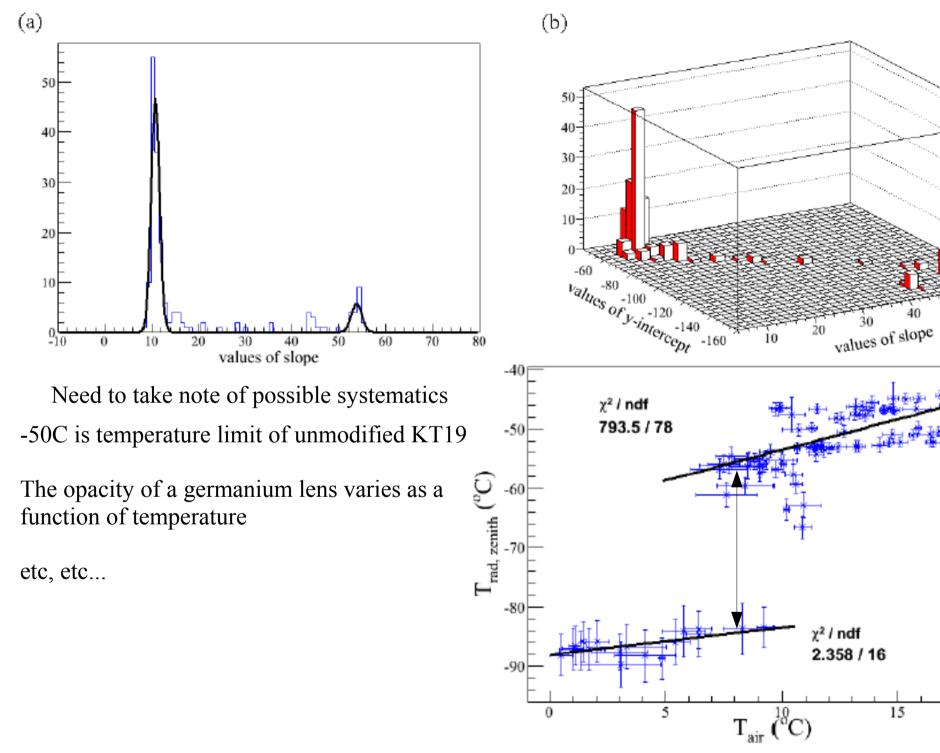
Hazy sky:

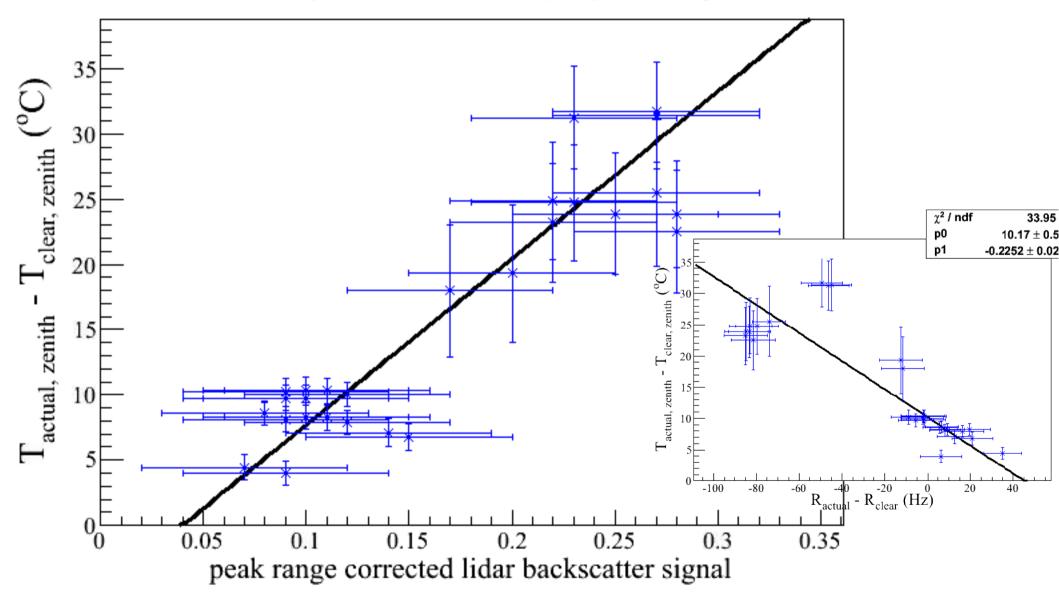
no obvious clouds

- at least 7 runs with $>35^{\circ}$ spread in zenith angles
- full 28 minute observing run (no hardware, or sun/moon rise contamination)
- lidar observes aerosol layer up to 3.5km

Leads to 28 observing runs from July->March







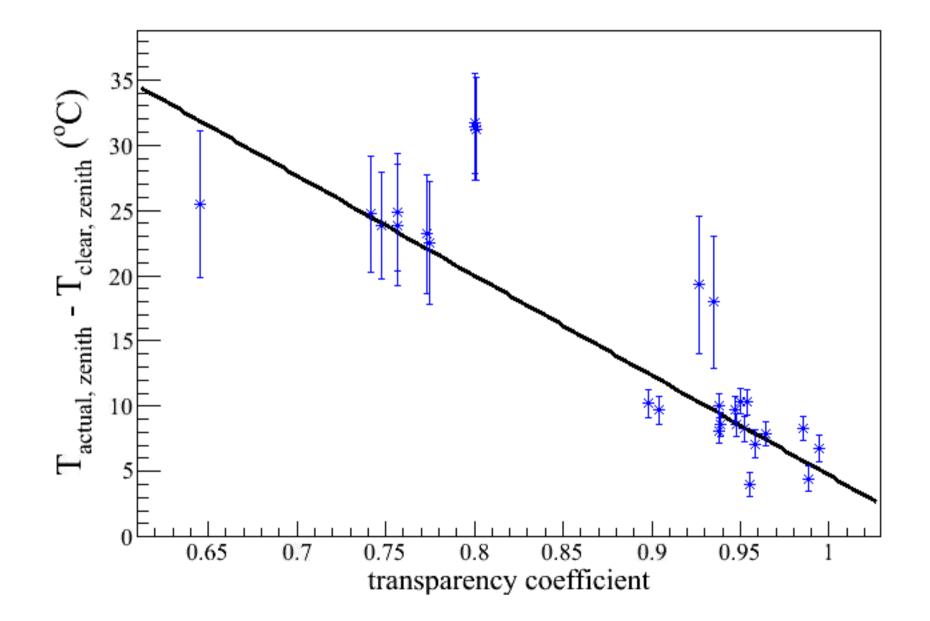
Observe a strong correlation between sky brightness temperature and aerosols

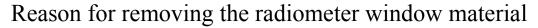
Summary

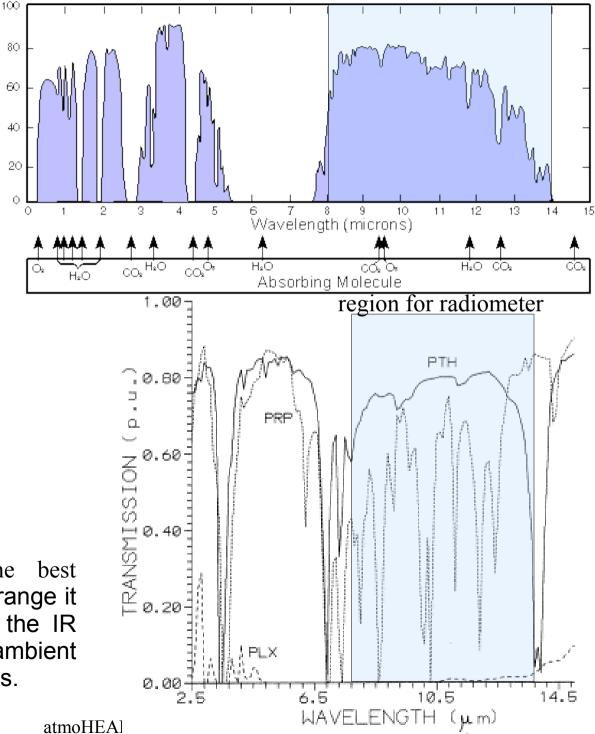
- •Aerosols can contribute a measurable amount to the equivalent sky brightness temperature between 8-14µm.
- •An IR radiometer is a useful addition to the suite of atmospheric monitoring tools
- •Measuring the IR can give an estimate of sky clarity independent of telescope systematics; and as it is a passive measurement it will not interfere with an observation
- •Whilst not sensitive to the aerosol distribution the IR is sensitive to low altitude aerosols in the crossover region a lidar is not
- •Measurements can be made during daytime to give a forecast of the night's observing conditions

extra slides:

- * IR brightness temperature versus transparency co-efficient
- * Radiometer window contribution to IR flux & brightness temperature
- * Temperature lapse rate
- * extra nights for 2011
- * initial study in 2009 with Leosphere lidar
- * Luminosity versus temperature curve for KT19 with Germanium lens
- * Lidar overlap factor







IR transmission of atmosphere

Transmittance (percent)

IR transmission of polyethylene (PTH)

Whilst polythene has the best window in the 8 to $<14\mu$ m range it will contribute greatly to the IR flux, biasing it toward ambient ground temperature values.

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Temperature Lapse Rate Γ_a

Taking the atmosphere to be in hydrostatic equilibrium, transparent to all radiation and containing no liquid particles, the first law of thermodynamics gives

$$C_V dT + P dV = dq = 0$$

Differentiate the equation of state, taking $\rho = 1/V$

$$PdV + VdP = \frac{k}{MM_0}dT$$

For an ideal gas

$$\frac{k}{MM_0} = C_P - C_V$$

Combining and replacing for $dP/dz = -g\rho$ gives

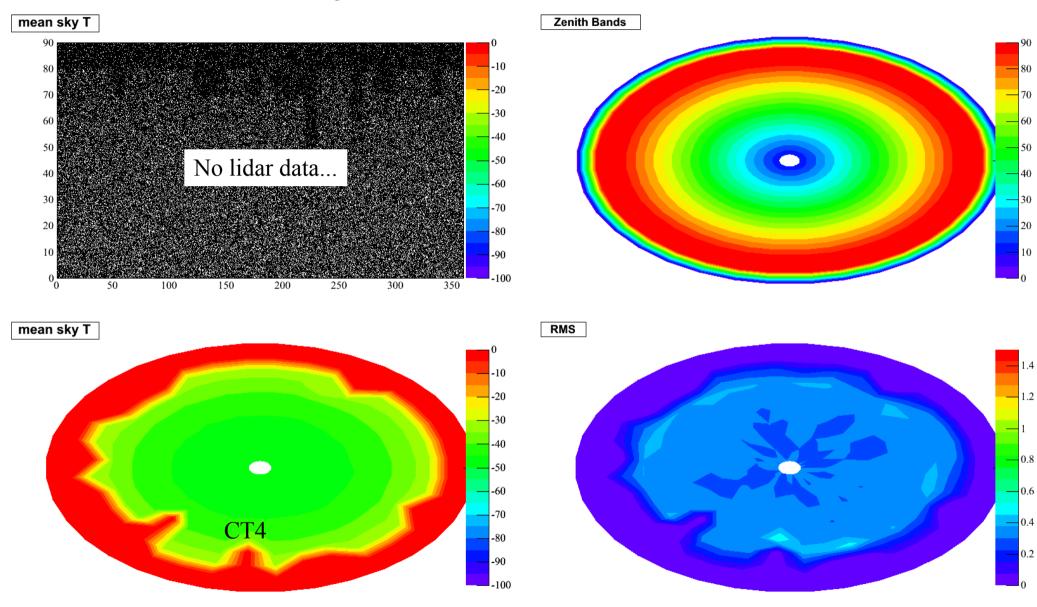
$$\frac{dT}{dz} = \frac{-g}{C_p} = -\Gamma_a$$

 $T(z) = T_0 - \Gamma_a z$

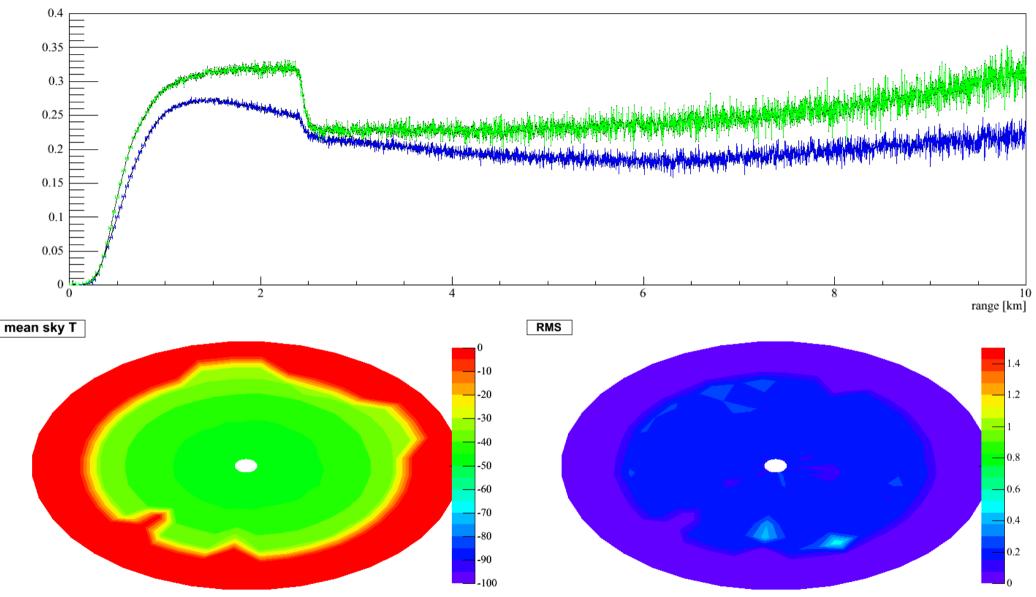
For a dry atmosphere $\Gamma_a \sim 9$ K/km The latent heat released by water vapour condensing out of the air (RH>60%) serves to raise this to ~6.5K/km

nb, all these nights listed as clear in shift logs & web summary.

"normal" observing conditions 2011/08/31



nb, all these nights listed as clear in shift logs & web summary.

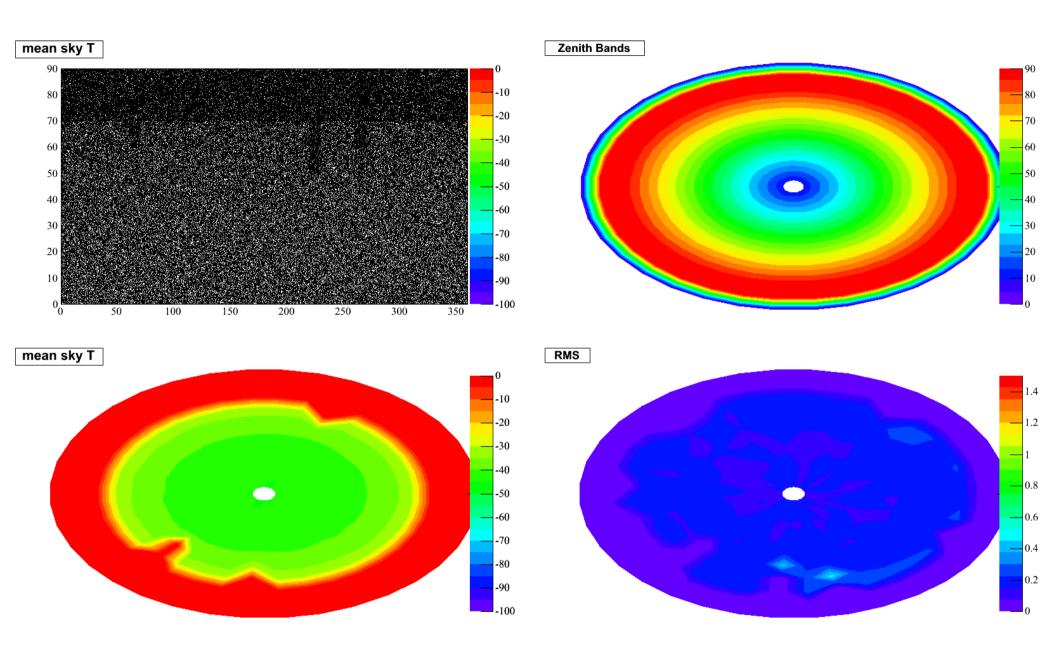


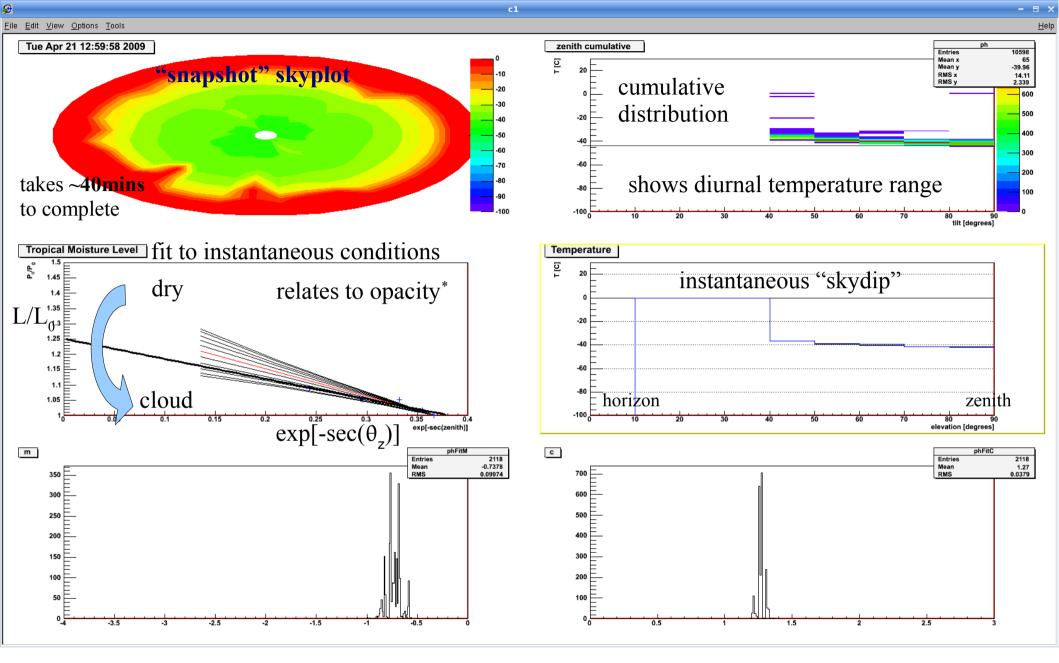
"normal" observing conditions 2011/08/30

3 tel runs, can't compare rates.

23

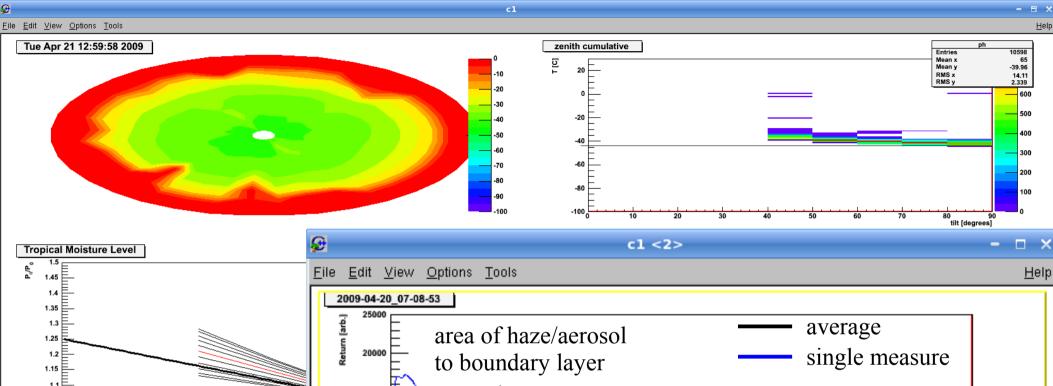
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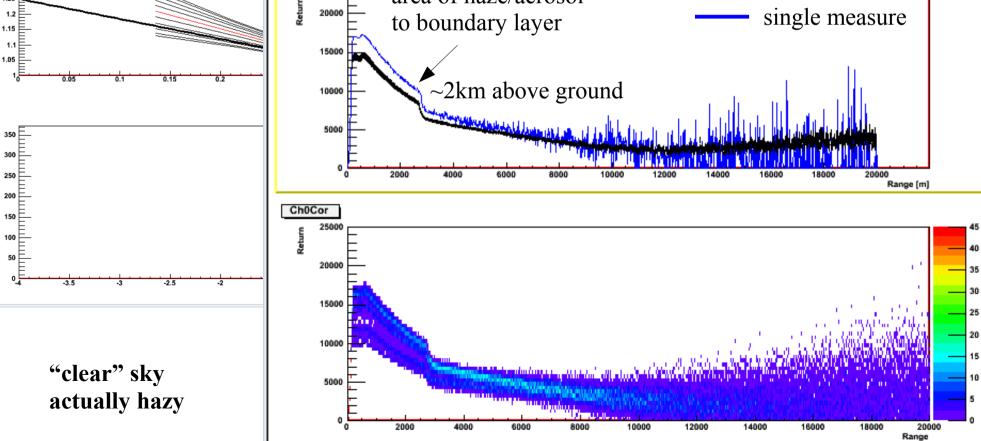
* Daniel, PhD Thesis, University of Durham (2002). Unsworth & Dalrymple Q J Roy Met Soc (1976).

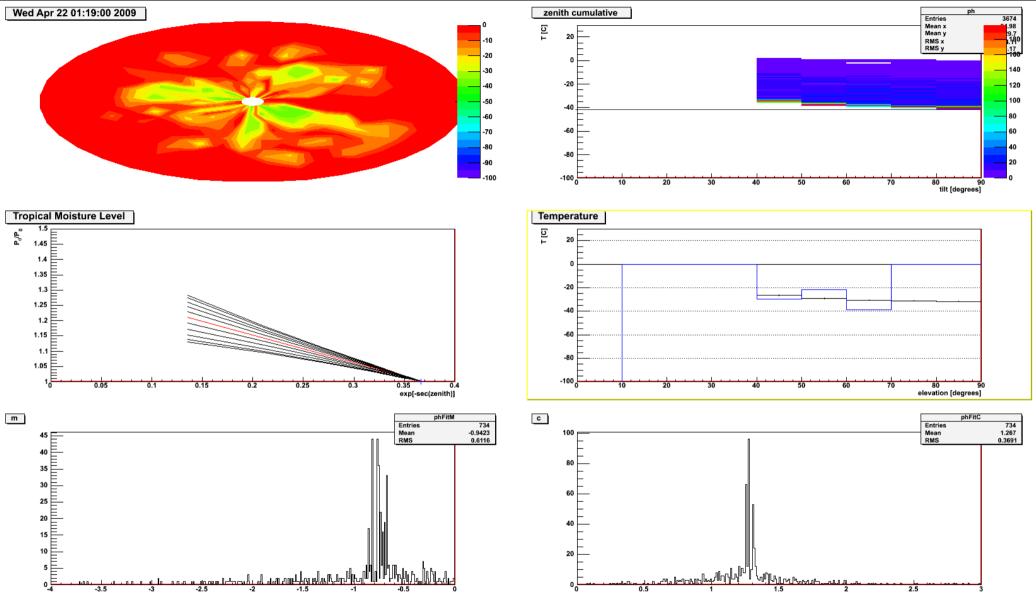
"clear" sky



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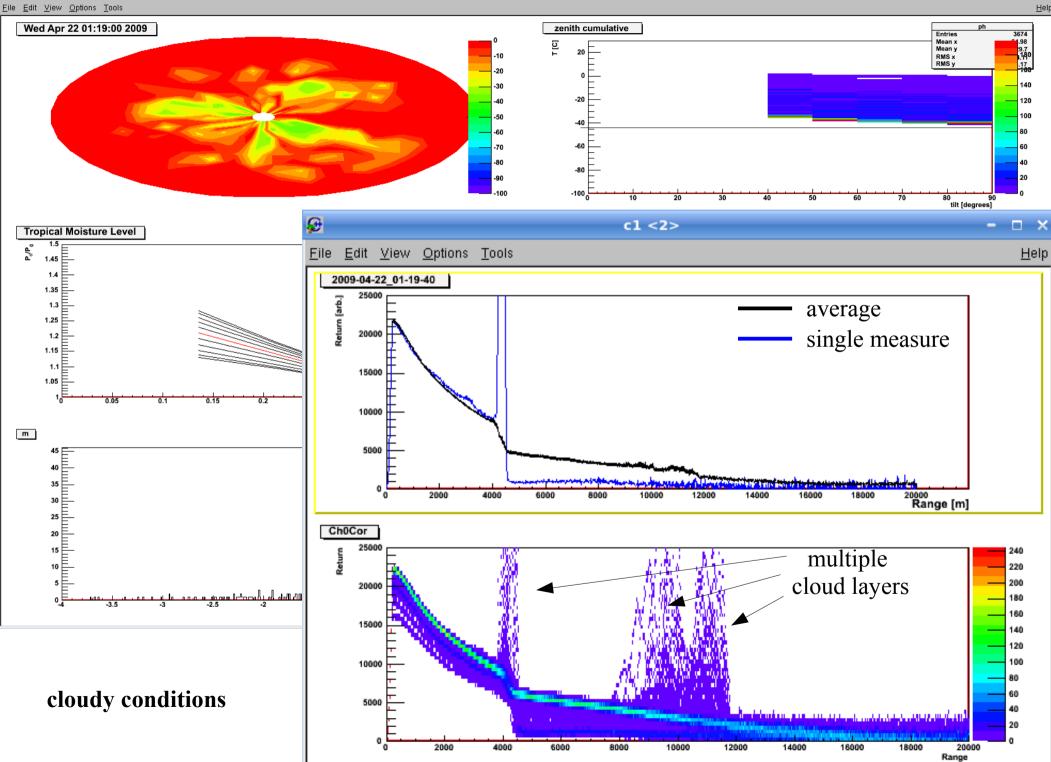


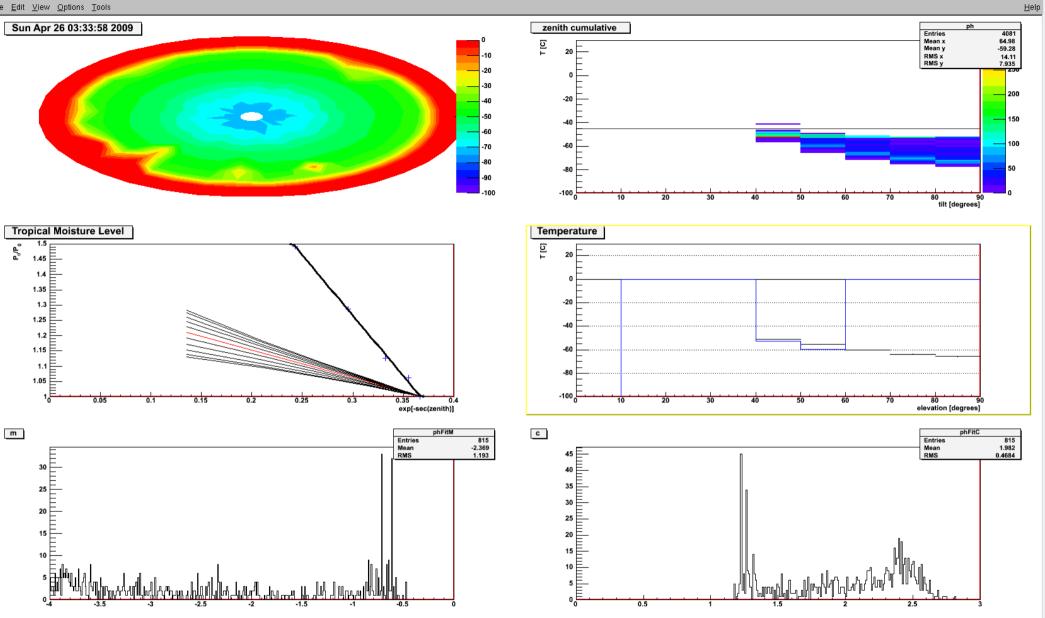


c1

cloudy conditions

- ⊟ × <u>H</u>elp

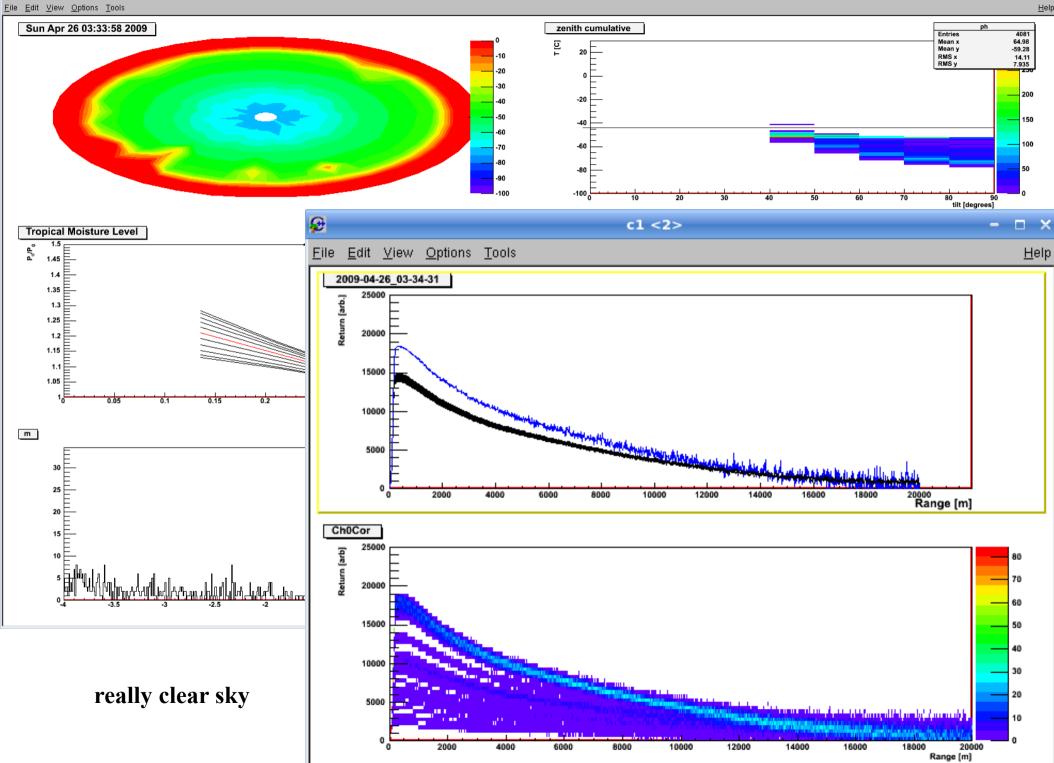




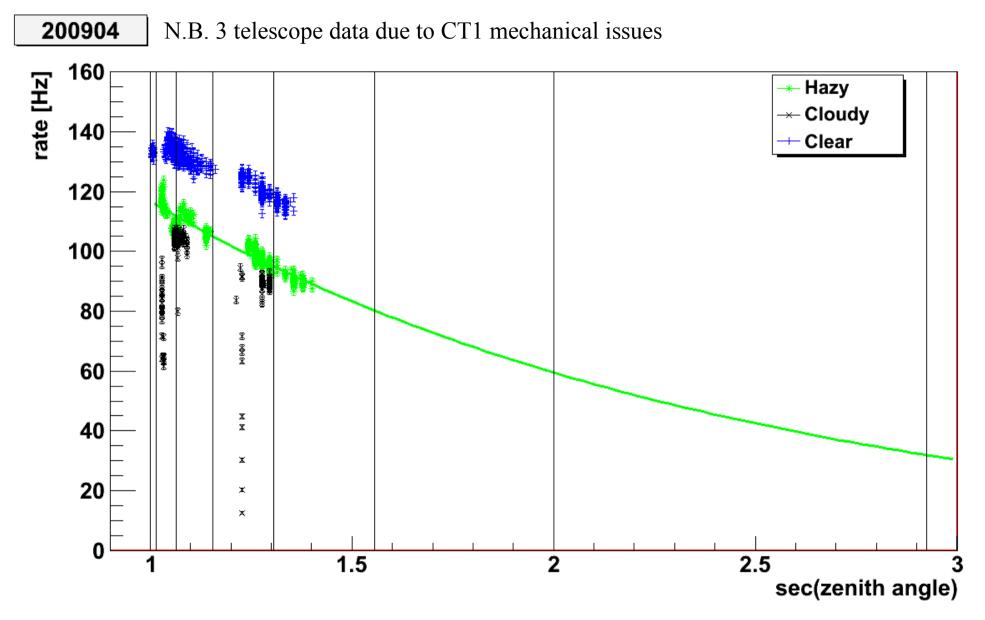
c1

really clear sky

- 🗉 🗙

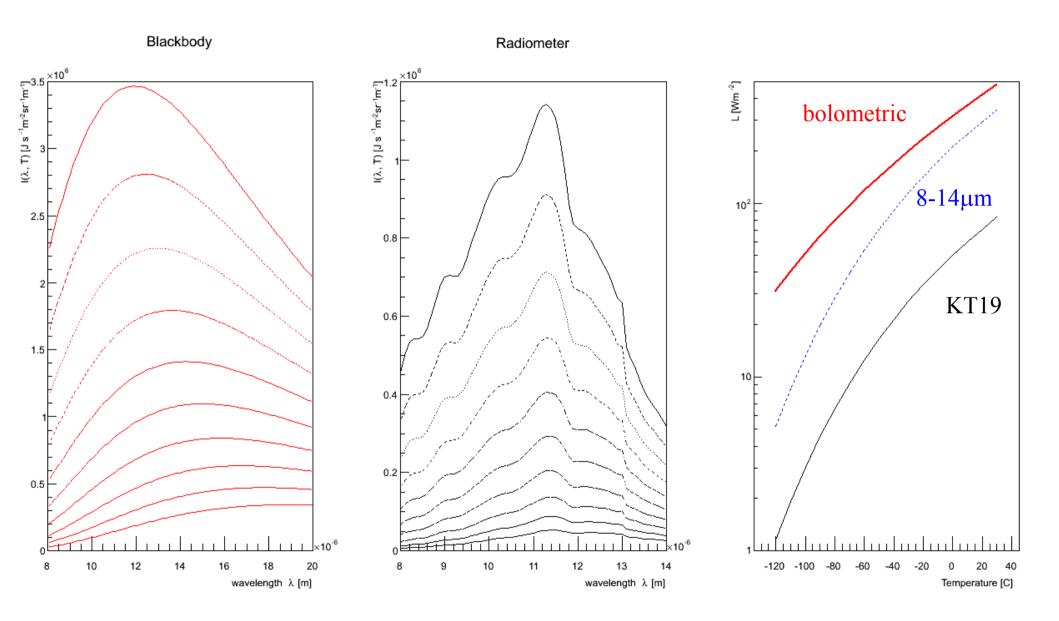


Trigger rates as a function of zenith angle θ for those nights in April 2009...



All 3 nights have similar temperatures & relative humidities

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Lidar overlap factor

