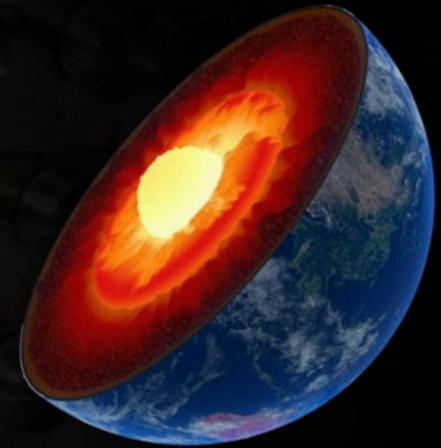


# Revealing the Earth's interior with geoneutrinos



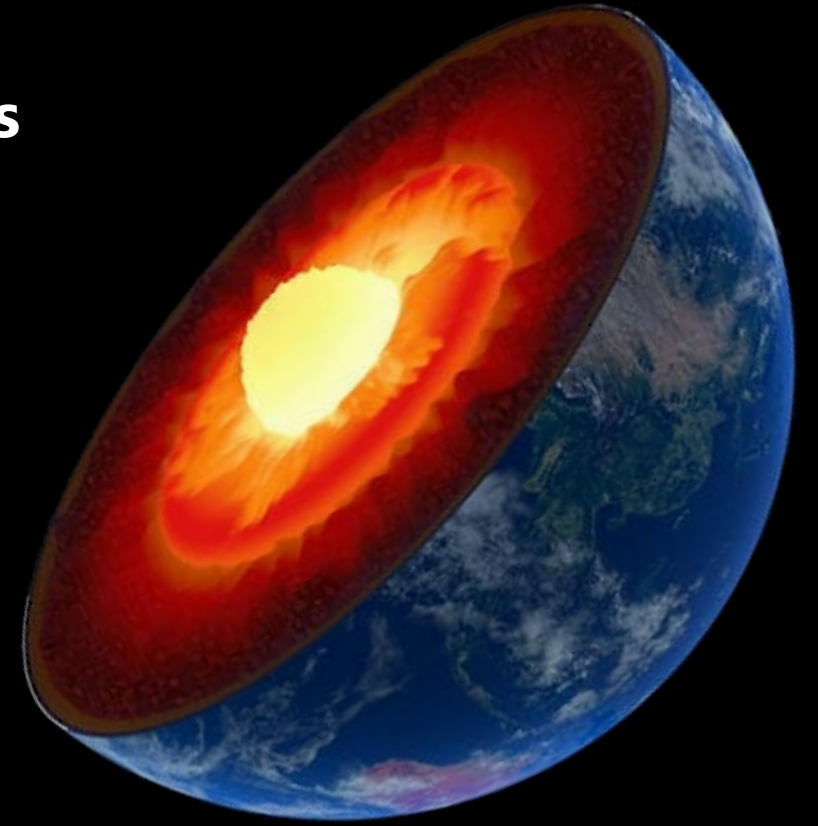
**Virginia Strati**

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Multi-messenger Tomography of the Earth 2023  
4-7 July 2023 - Paris

# Outline

- **Natural radioactivity and Heat power in the Earth:  
open questions**  
**Virginia Strati**
- **Overview of experimental results on geoneutrinos**  
**Livia Ludhova**
- **What can we learn from geoneutrinos results?**  
**Virginia Strati**
- **Perspectives and challenges for  
geoneutrino experiments**  
**Fabio Mantovani**
- **JUNO's perspectives for geoneutrinos**  
**Yury Malyshkin**





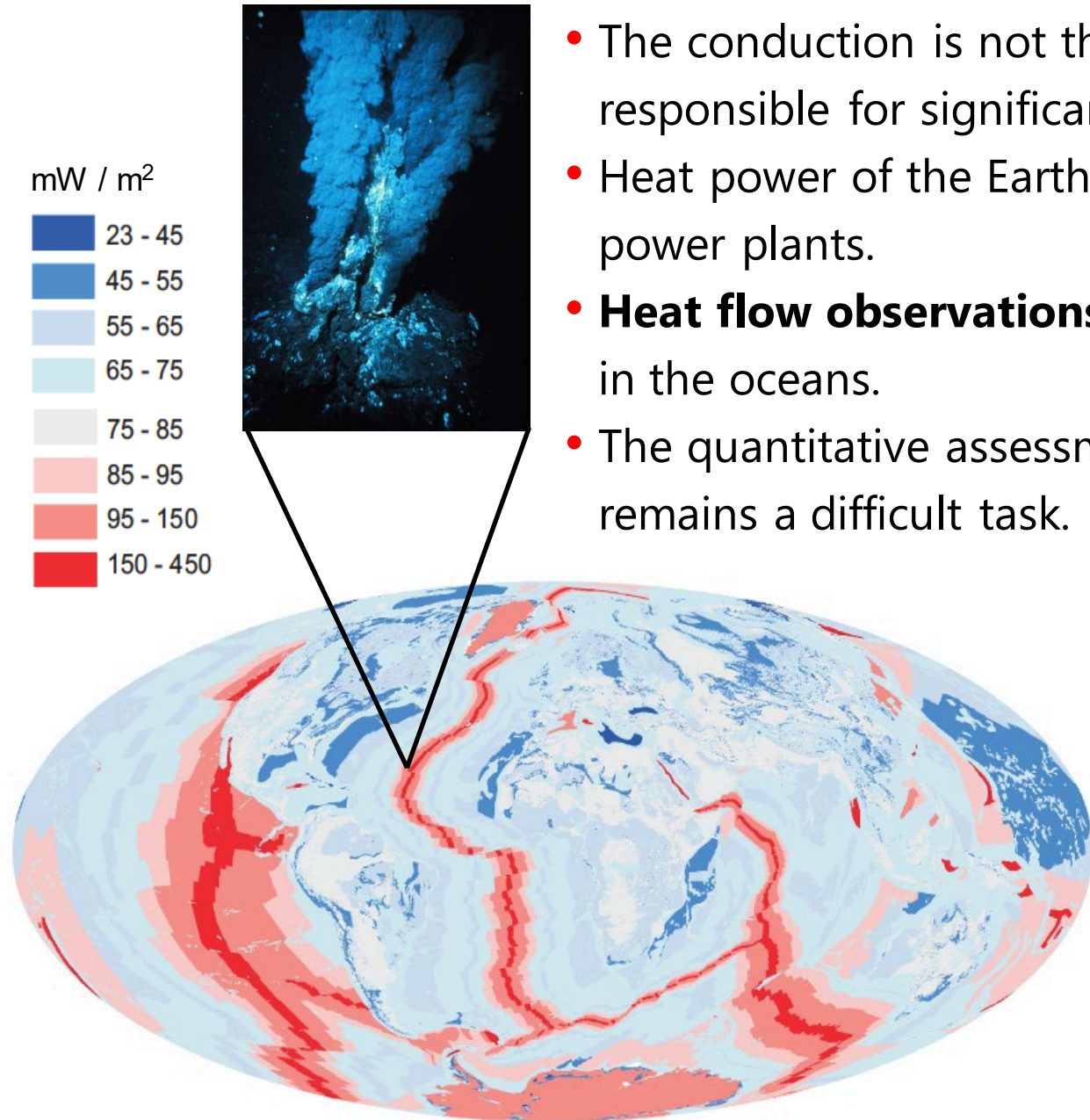
# Open questions about natural radioactivity in the Earth



- **What is the radiogenic contribution to terrestrial heat production?**
- **At which thermal conditions the Earth initially is formed?**
- **How much U and Th are in the crust and in the mantle?**
- **What is hidden in the Earth's core?**



# Heat power of the Earth



- The conduction is not the only way of Earth's cooling: **convective motions** are responsible for significant fraction of surface heat loss.
- Heat power of the Earth **Q [30-49 TW]** is the equivalent of  $\sim 10^4$  nuclear power plants.
- **Heat flow observations** are sparse, non-uniformly distributed and not reliable in the oceans.
- The quantitative assessment of heat transport by **hydrothermal circulation** remains a difficult task.

REFERENCE	Continents	Oceans	Total
	q <sub>CT</sub> [mW m <sup>-2</sup> ]	q <sub>OCS</sub> [mW m <sup>-2</sup> ]	Q (TW)
Williams et al., 1974	61	92	43 ± 6
Davies, 1980	55	95 ± 10	41 ± 4
Sclater et al., 1980	57	99	42
Pollack et al., 1993	65 ± 2	101 ± 2	44 ± 1
Hofmeister and Criss, 2005	61	65	31 ± 1
Jaupart et al., 2015	65	107	46 ± 2
<b>Davies and Davies, 2010</b>	<b>71</b>	<b>105</b>	<b>47 ± 2</b>
Davies, 2013	65	96	45
Lucazeau, 2019	66.7	89.0	44

# The Earth's heat sources

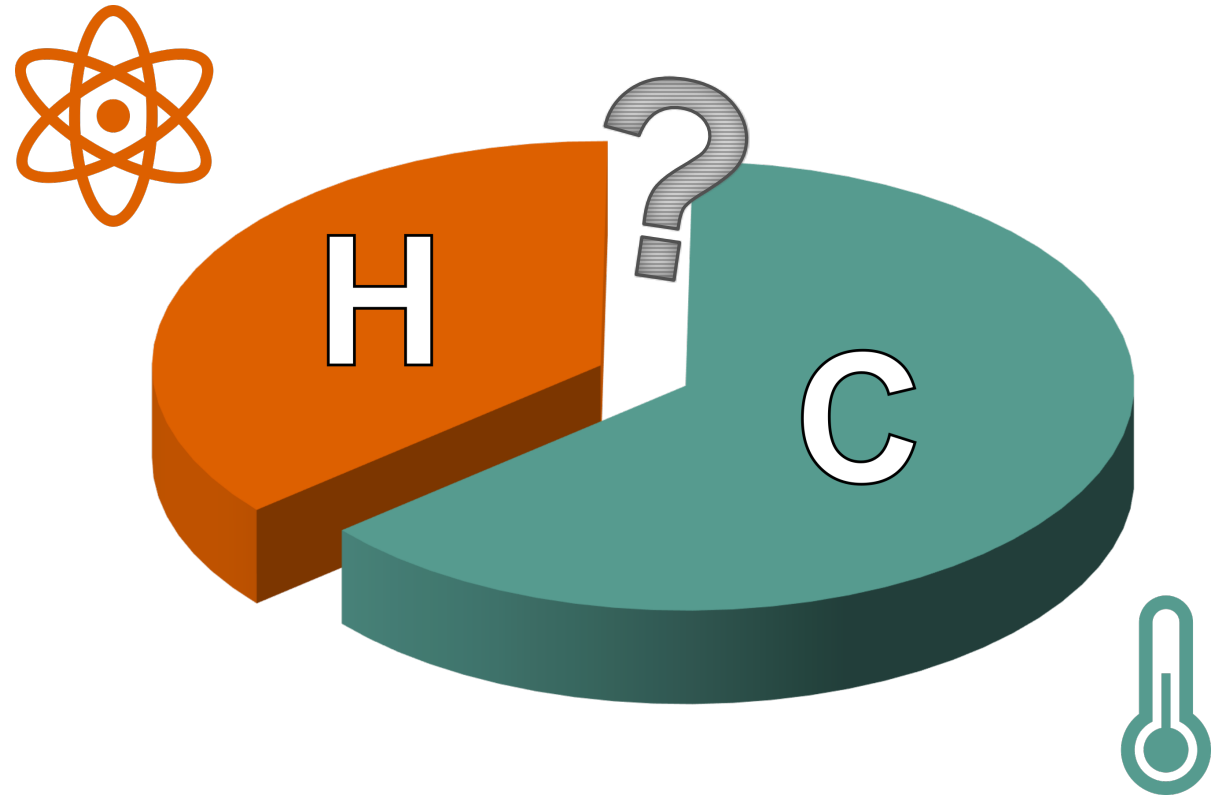
Neglecting tidal dissipation and gravitation contraction (<0.5 TW), the two contributions to the total heat loss ( $Q$ ) are:

- **Secular Cooling (C)**: cooling down caused by the initial hot environment of early formation's stages
- **Radiogenic Heat (H)** due to naturally occurring decays of Heat Producing Elements (HPEs), i.e. U, Th and K, inside our planet.
- The Urey ratio is a key parameter that can be seen as the ratio of **H** over **Q**. It measures the efficiency of the Earth's convective engine in evacuating heat generated by radioactive decays.

$$U_R = \frac{H}{Q} \text{ (Bulk Earth)}$$

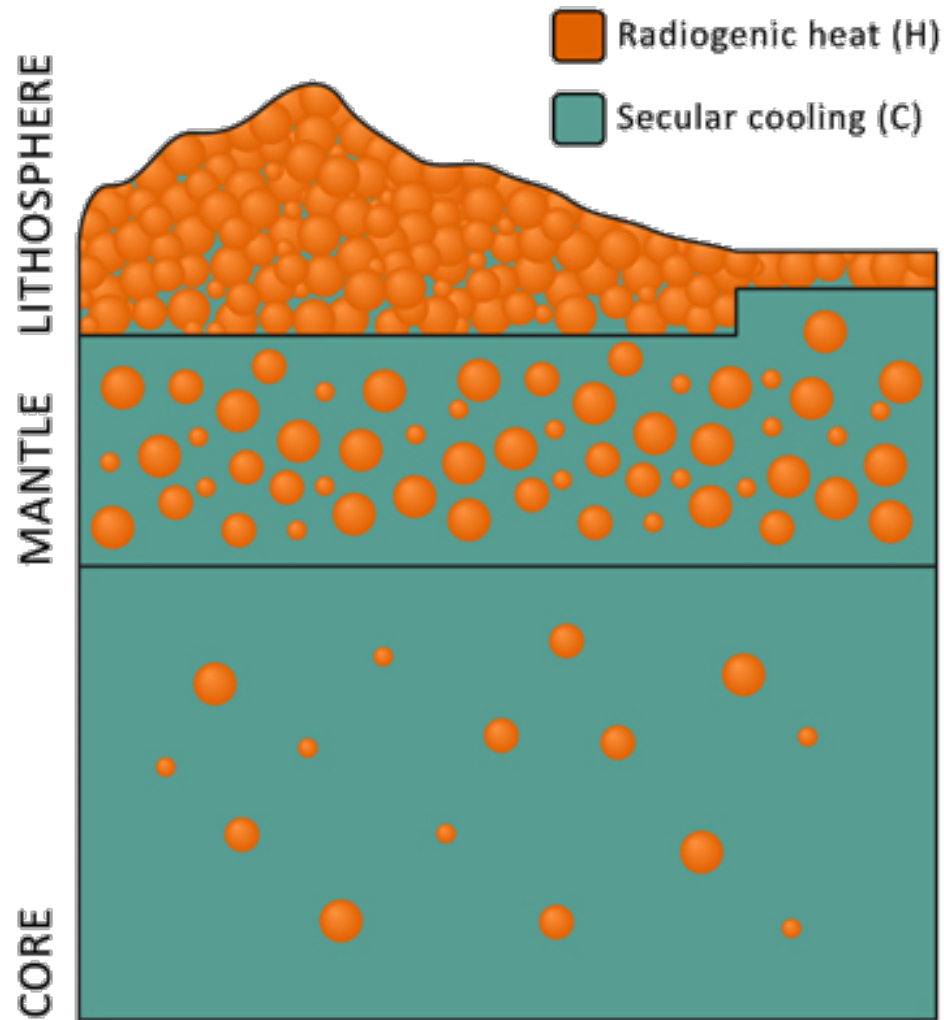
$$U_R = \frac{H - H_{CC}}{Q - H_{CC}} \text{ (Convective)}$$

$H_{CC}$  : radiogenic power of the continental crust





# The Earth's budget



- $H_{CC}$  = radiogenic power of the continental crust
- $H_{OC}$  = radiogenic power of the oceanic crust
- $H_{CLM}$  = radiogenic power of the continental lithospheric mantle

$$C = Q - H$$

$$C_M = Q - H - C_C$$

$$H_M = H - H_{LS} - H_C$$

$$H_{LS} = H_{CC} + H_{OC} + H_{CLM}$$

$$U_R = \frac{H - H_{CC}}{Q - H_{CC}}$$

	Range [TW]	Adopted [TW]
H	[10 ; 37]	$19.3 \pm 2.9$
$H_{LS}$	[6 ; 11]	$8.1^{+1.9}_{-1.4}$
$H_M$	[0 ; 31]	$11.0^{+3.3}_{-3.4}$
$H_C$	[0 ; 5]	0

	Range [TW]	Adopted [TW]
C	[8 ; 39]	$28 \pm 4$
$C_{LS}$	~ 0	0
$C_M$	[1 ; 29]	$17 \pm 4$
$C_C$	[5 ; 17]	$11 \pm 2$

- The mass of the **lithosphere** (~ 2% of the Earth's mass) contains ~ 40% of the total HPEs:  **$H_{LS} \sim 8$  TW**.
- Radiogenic power of the **mantle**  $H_M$  and the contributions to C from mantle ( $C_M$ ) and core ( $C_C$ ) are model dependent.

# The building blocks of the Bulk Silicate Earth

- The **Bulk Silicate Earth (BSE)** describes the primordial non-metallic Earth condition that followed planetary accretion and core separation, prior to its differentiation into a mantle and crust.
- We can tempt to naively build the BSE in the image and likeness of a primitive meteorite: every year  $10^4$  kg of meteorites falling to the ground
- The **chondrites** are undifferentiated rocky meteorites: they represents the initial «cocktail» of the Solar System

## **CARBONACEOUS (CI):**

composition close to Solar photosphere  
- high concentration of oxides



## **ENSTATITE (EH) :**


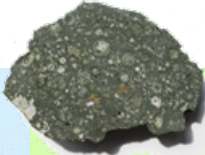
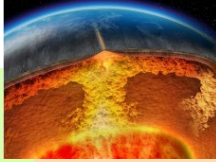
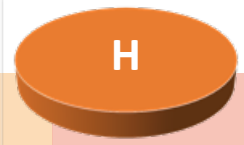
Isotopic composition more similar to the terrestrial samples - high metallic iron content



References	Chondrite	$a_{\text{BSE(U)}} [\text{ng g}^{-1}]$
Jackson and Jellinek [160]	CI-EH	$14 \pm 3$
O'Neill and Palme [67]	CI-EH	10
Javoy and Kaminski [157]	EH	$15 \pm 2$
Javoy et al. [101]	EH	$12 \pm 2$
McDonough and Sun [38]	C	$20 \pm 4$
Lyubetskaya and Korenaga [161]	C	$17 \pm 3$
Palme and O'Neill 2007	C	$22 \pm 3$
Arevalo 2010	C	$20 \pm 4$
Wang et al. 2018	C	$20 \pm 2$
Palme and O'Neill [162]	C	$23 \pm 3$
Turcotte 2002*	CI-EH	$35 \pm 4$
Turcotte 2014	CI-EH	31



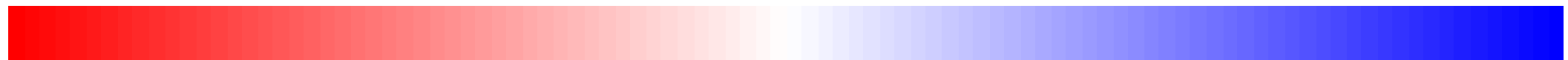
# Different models of Bulk Silicate Earth

	 <b>CC</b>	 <b>GC</b>	 <b>GD</b>	 <b>FR</b>
	<b>COSMOCHEMICAL</b> <ul style="list-style-type: none"> <li>• Enstatites chondrites</li> <li>• Sufficiently high iron content to explain the metallic core</li> </ul>	<b>GEOCHEMICAL</b> <ul style="list-style-type: none"> <li>• Chondritic compositions for refractory lithophile elements</li> <li>• Constraints from terrestrial samples</li> </ul>	<b>GEODYNAMICAL</b> <ul style="list-style-type: none"> <li>• Energetics arguments of mantle convection</li> <li>• Observed surface heat loss</li> </ul>	<b>FULLY RADIOGENIC</b> <ul style="list-style-type: none"> <li>• The terrestrial heat (47 TW) is assumed to be fully accounted by radiogenic production</li> </ul>
H (U+Th+ K) [TW]	<b>11</b>	<b>20</b>	<b>34</b>	<b>47</b>
M (U) [ $10^{16}$ kg]	5	8	14	20

Temperature of the Earth at initial stage of its formation



+





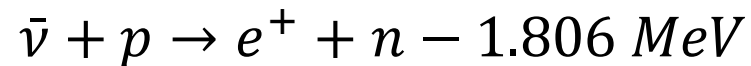
# Geo-neutrinos: anti-neutrinos from the Earth

U, Th and  $^{40}\text{K}$  in the Earth release heat together with anti-neutrinos, in a **well-fixed ratio**:

Decay	$T_{1/2}$ [ $10^9$ yr]	$E_{\text{max}}$ [MeV]	$Q$ [MeV]	$\varepsilon_{\bar{\nu}}$ [ $\text{kg}^{-1}\text{s}^{-1}$ ]	$\varepsilon_H$ [W/kg]
$^{238}\text{U} \rightarrow ^{206}\text{Pb} + 8\ ^4\text{He} + 6e + 6\bar{\nu}$	4.47	3.26	51.7	$7.46 \times 10^7$	$0.95 \times 10^{-4}$
$^{232}\text{Th} \rightarrow ^{208}\text{Pb} + 6\ ^4\text{He} + 4e + 4\bar{\nu}$	14.0	2.25	42.7	$1.62 \times 10^7$	$0.27 \times 10^{-4}$
$^{40}\text{K} \rightarrow ^{40}\text{Ca} + e + \bar{\nu}$ (89%)	1.28	1.311	1.311	$2.32 \times 10^8$	$0.22 \times 10^{-4}$

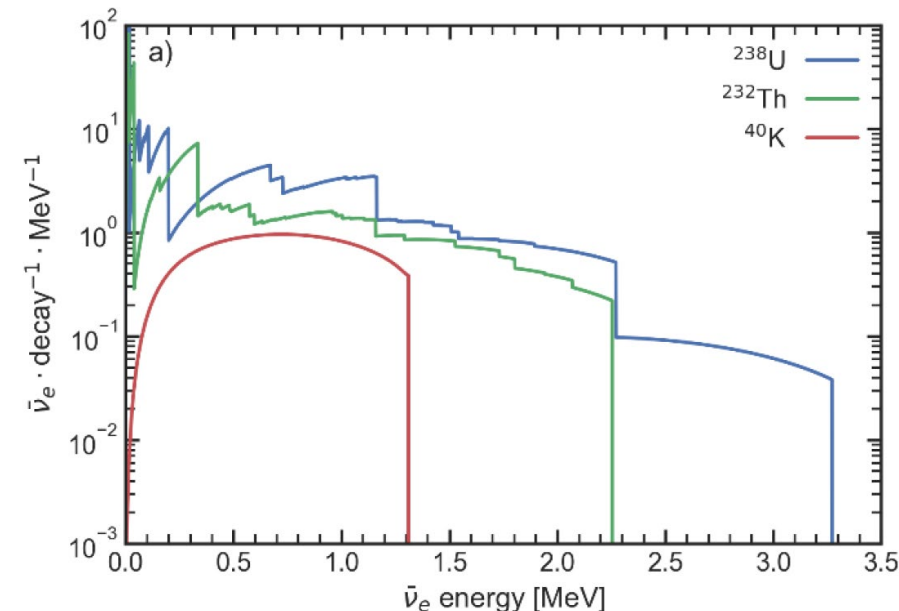
- Earth emits (mainly) antineutrinos  $\Phi_{\bar{\nu}} \sim 10^6 \text{ cm}^{-2}\text{s}^{-1}$  whereas Sun shines in neutrinos

- A fraction of geo-neutrinos from U and Th (not from  $^{40}\text{K}$ ) are above threshold for inverse  $\beta$  on protons:



- Different components can be distinguished due to different energy spectra: e. g. anti- $\nu$  with highest energy are from U

- Signal unit: **1 TNU** = one event per  $10^{32}$  free protons/year



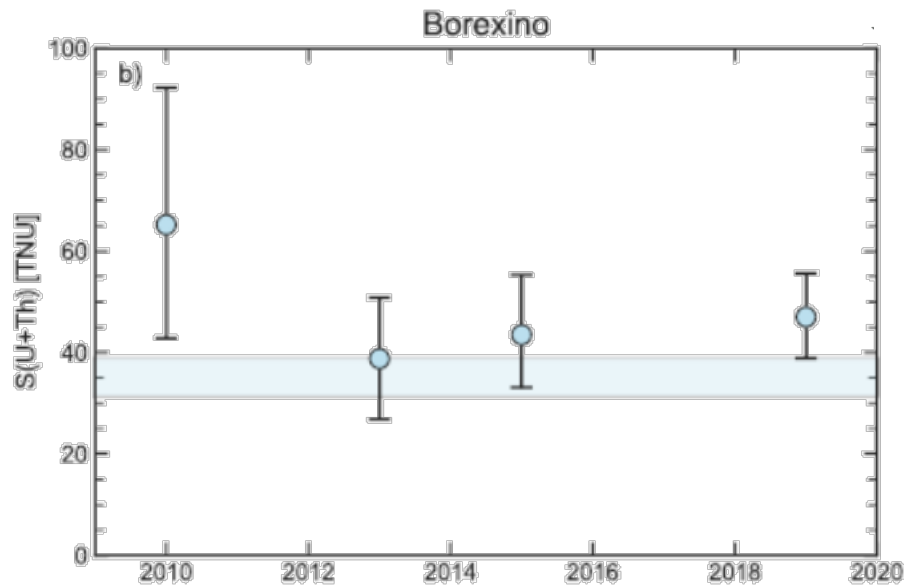




# Borexino geoneutrino results

## Borexino

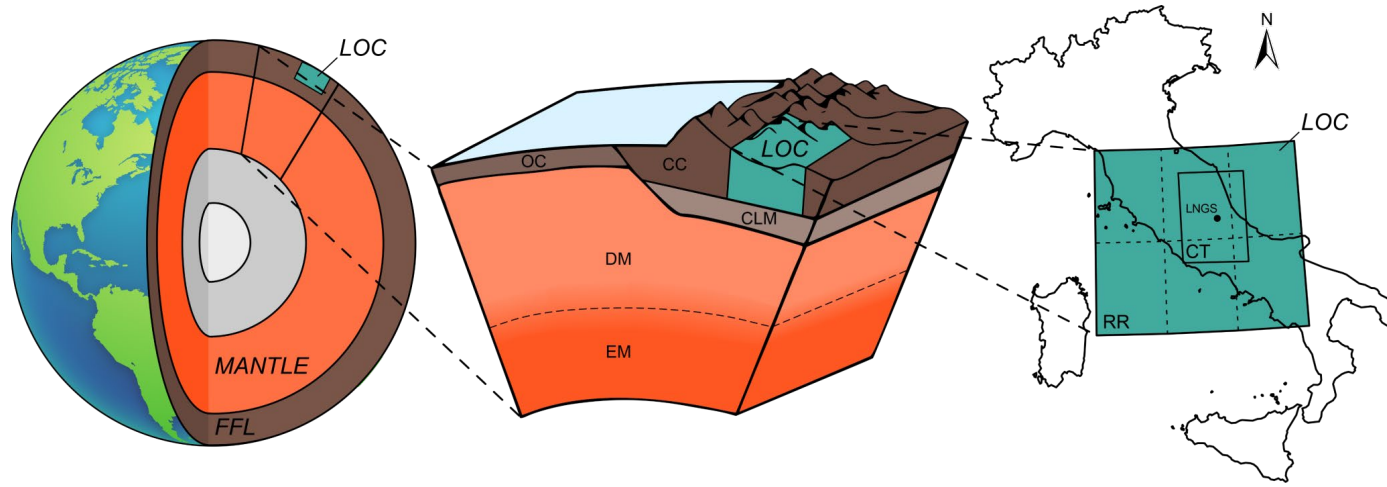
- Period: 2007 – 2019
- Geo- $\nu$  events:  $52.6^{+7.4}_{-6.3}$
- Signal:  $47.0^{+8.7}_{-7.9}$  TNU



- Horizontal bars traces the **expected signal** at  $1\sigma$  C.L.
- In the second decade of the 21<sup>st</sup> century the results published with greater statistical significance highlighted the necessity of **geophysical and geological models** for understanding geoneutrino signal.

# Mantle geoneutrino signals from experimental signal

$$S_{Exp}^i(U + Th) = S_M^i(U + Th) + S_{FFL}^i(U + Th) + S_{LOC}^i(U + Th)$$



- U and Th distributed in the **Local Crust (LOC)** (i.e.  $\sim 500$  km within the detector) gives a significant contribution to the signal ( $\sim 50\%$  of the total).

$$S_M^i(U + Th) = S_{Exp}^i(U + Th) - S_{FFL}^i(U + Th) - S_{LOC}^i(U + Th)$$

- The signal of the **Far Field Lithosphere (FFL)** is modeling based on global reference model.
- The **Local Crust (LOC)** modeling should be built with geochemical and/or geophysical information typical of the local regions.

$S_{LOC}(U + Th)$	$9.2 \pm 1.2$ TNU
$S_{FFL}(U + Th)$	$16.3^{+4.8}_{-3.7}$ TNU



# Measurements vs models

$$S_{\text{Exp}} (\text{U+Th}) = 47.0^{+8.6}_{-8.1} \text{ TNU}$$

## BSE models according to different authors:

**J** = M. Javoy et al., EPSL 293, (2010).

**L&K** = T. Lyubetskaya and J. Korenaga, J. Geoph. Res. Sol. Earth, 112 (2007)

**T** = S. Taylor, Proc. Lunar Planet. Sci. Conf. 11, 333 (1980)

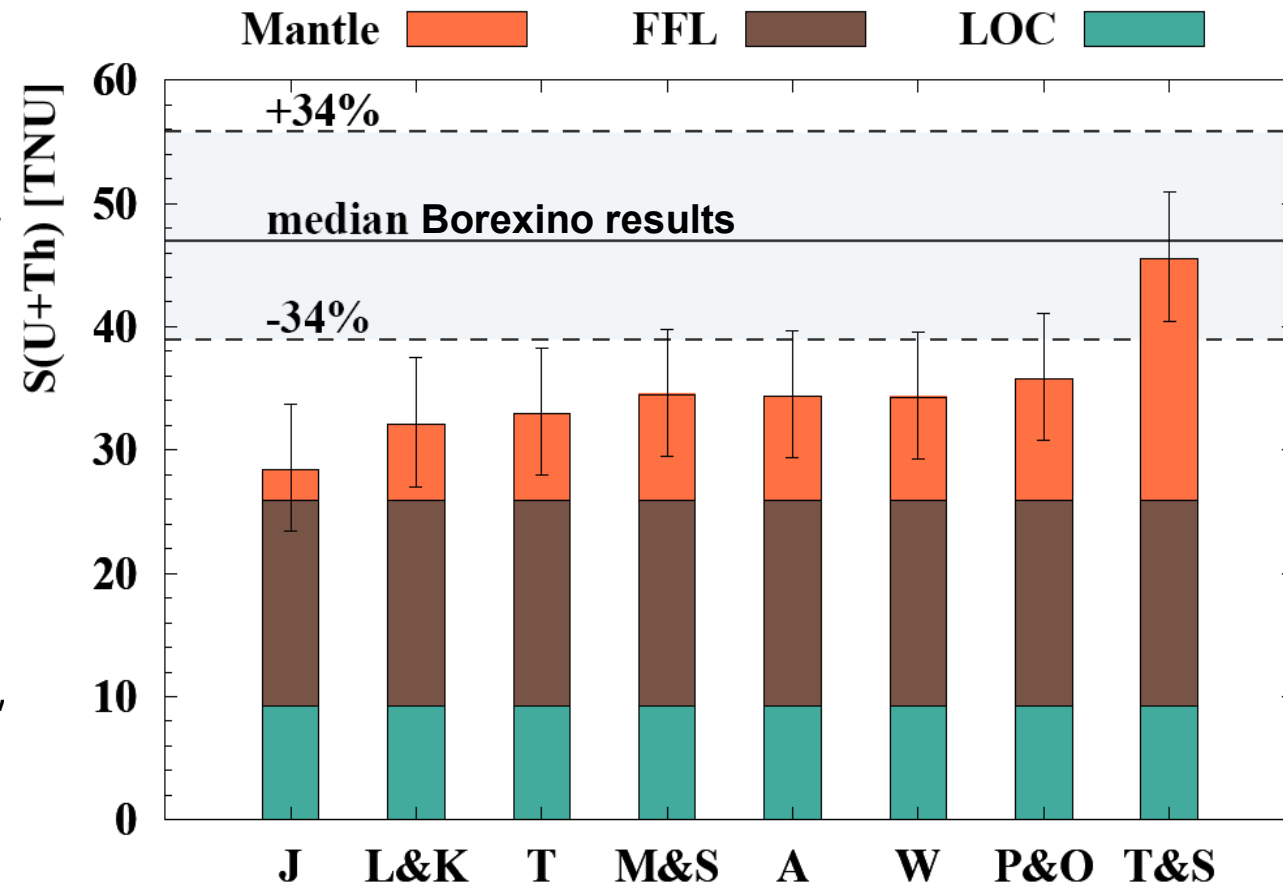
**M&S** = W. F. McDonough and S. Sun, Chem. Geol. 120, (1995)

**A** = D. L. Anderson, Cambridge University Press, (2007)

**W** = H. S. Wang et al., Icarus 299, (2018)

**P&O** = H. Palme and H. O'Neill, Treatise of Geochemistry, (2003)

**T&S** = D. L. Turcotte and G. Schubert, Cambridge University Press, (2002)



The Borexino observations favor geological models that predict a relatively **high concentration of radioactive elements** in the mantle.

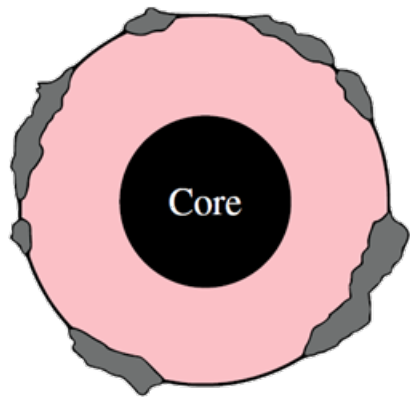
# Mantle radiogenic power (BX)

Since  $H_{LS} (U+Th) = 6.9_{-1.2}^{+1.6}$  TW is independent from the BSE model, the discrimination capability can be studied in the space  $S_M(U+Th)$  vs  $H_M(U+Th)$ :

$$S_M(U+Th) = \beta \cdot H_M(U+Th)$$

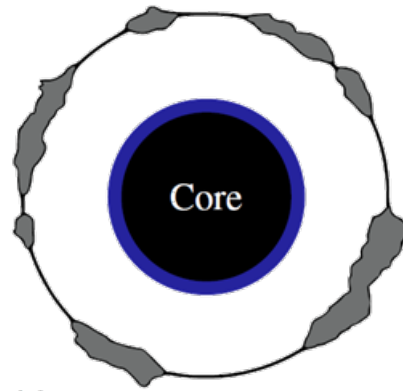
**High scenario**

$$\beta = 0.98 \text{ TNU/TW}$$



**Low scenario**

$$\beta = 0.75 \text{ TNU/TW}$$



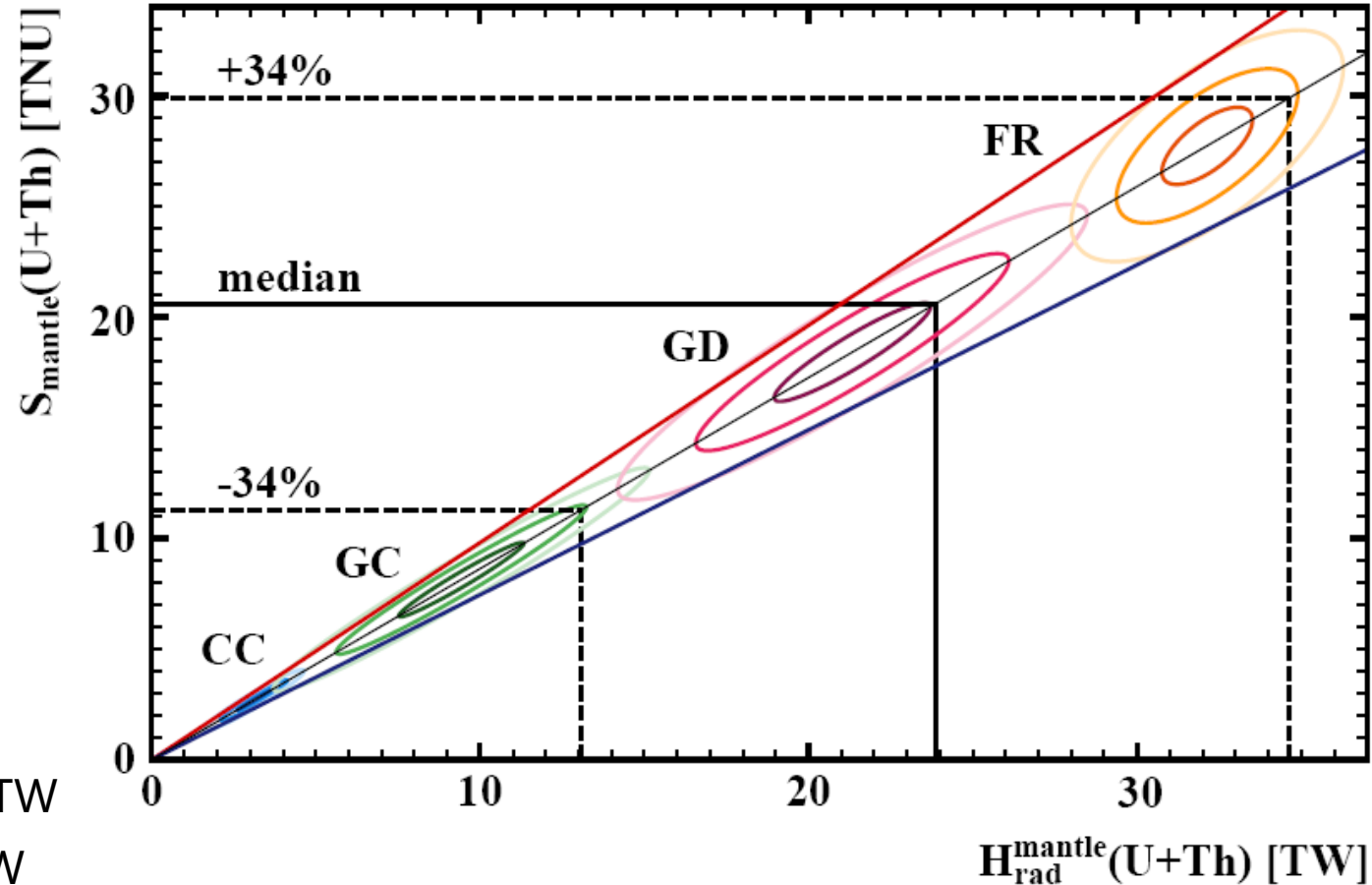
**Cosmochemical Model (CC)**  $H_M(U + Th) = 0.7 - 3.8$  TW

**Geochemical Model (GC)**  $H_M(U + Th) = 7.5 - 10.9$  TW

**Geodynamical Model (GD)**  $H_M(U + Th) = 19.8 - 23.3$  TW

**Fully radiogenic (FR)**  $H_M(U + Th) = 30.5 - 34.0$  TW

$$S_{\text{Mantle}} (U+Th) = 21.2_{-9.0}^{+9.6} \text{ TNU}$$



$$H_{\text{Mantle}} (U+Th) = 24.6_{-10.4}^{+11.1} \text{ TW}$$

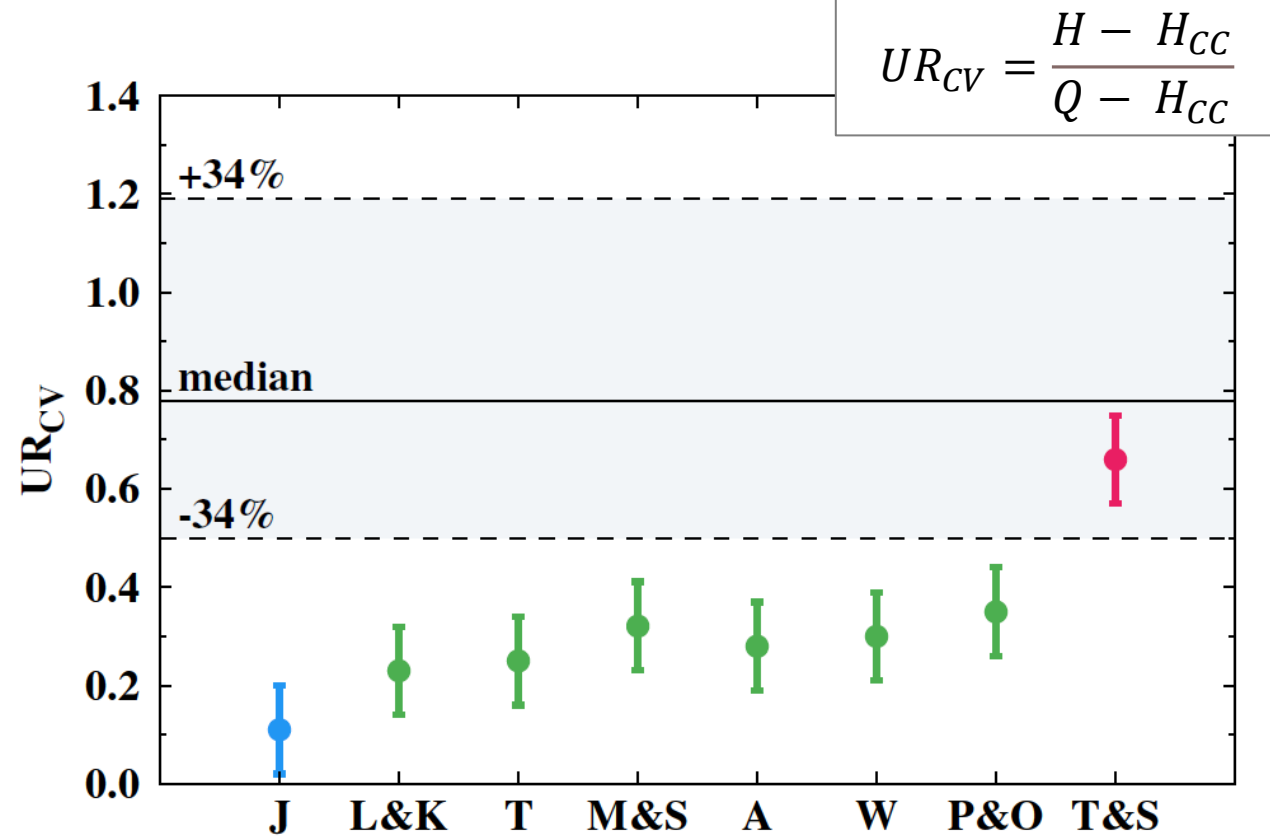
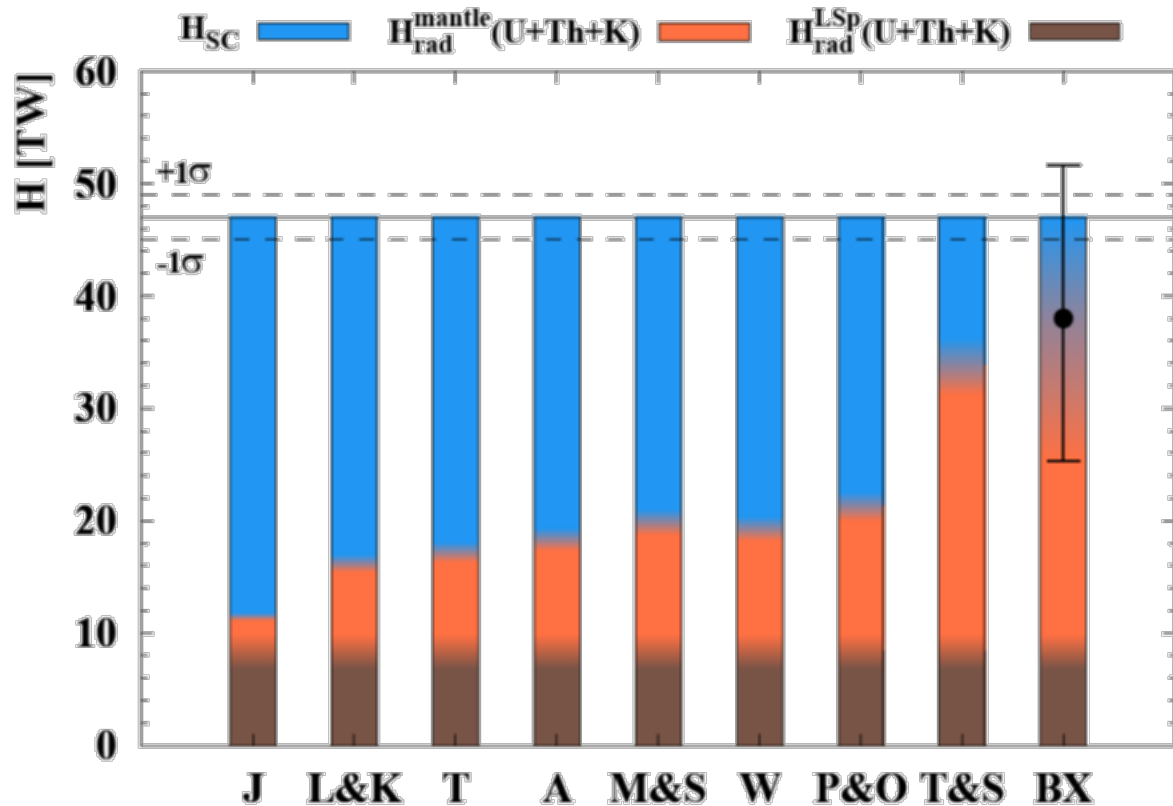
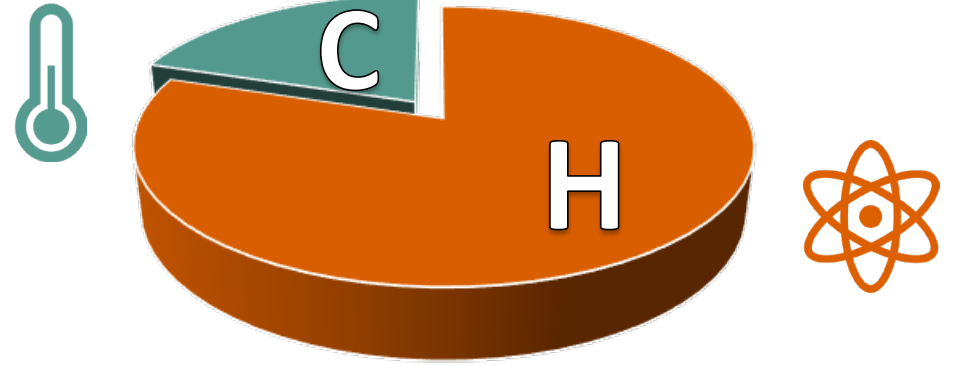


# Earth's heat budget (BX)

$$H_{LS} (U+Th+K) = 8.1_{-1.4}^{+1.9} \text{ TW}$$

Assuming the contribution from  $^{40}\text{K}$  to be 18% of the total mantle radiogenic heat

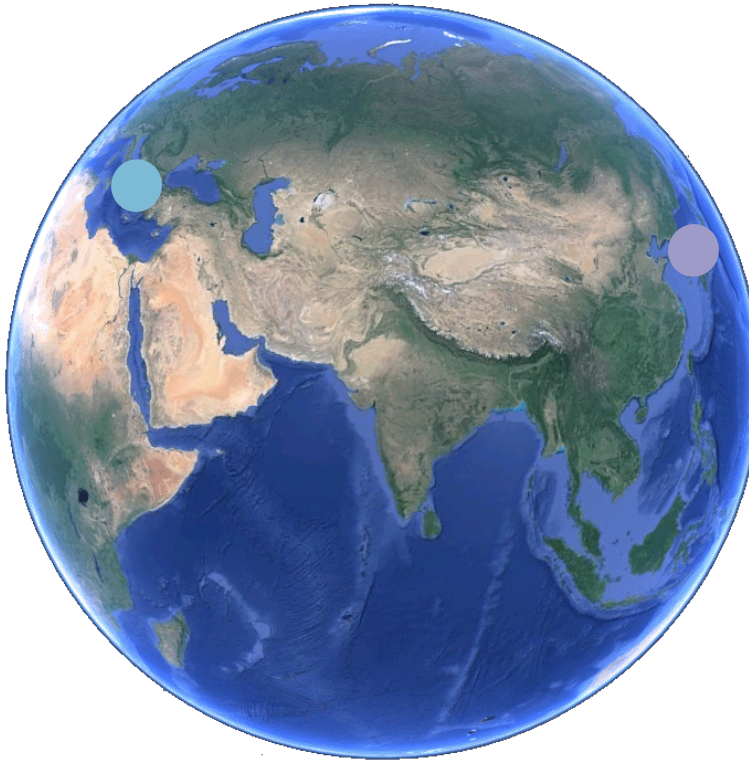
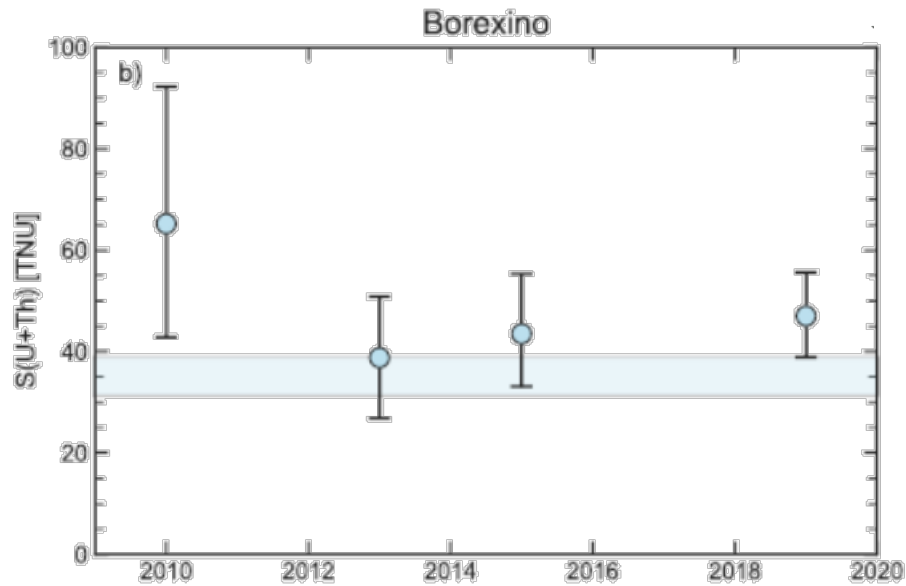
$$H_{\text{Earth}} (U+Th+K) = 38.2_{-12.7}^{+13.6} \text{ TW}$$



# Borexino and KamLAND geoneutrino results

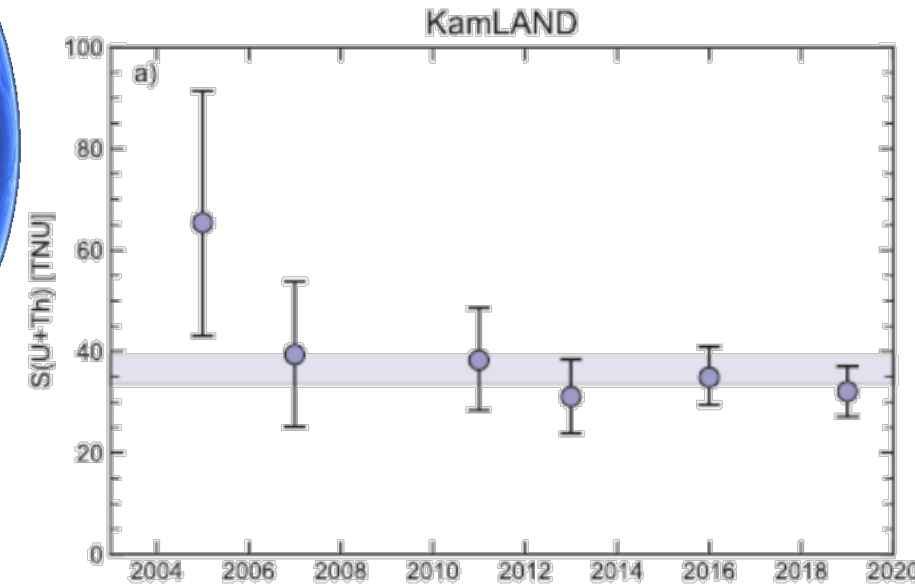
## Borexino

- Period: 2007 – 2019
- Geo- $\nu$  events:  $52.6^{+7.4}_{-6.3}$
- Signal:  $47.0^{+8.7}_{-7.9}$  TNU



## KamLAND

- Period: 2002 – 2019
- Geo- $\nu$  events:  $168.8^{+26.3}_{-26.5}$
- Signal:  $32 \pm 5$  TNU



- Horizontal bars traces the **expected signal** at  $1\sigma$  C.L.
- In the second decade of the 21<sup>st</sup> century the results published with greater statistical significance highlighted the necessity of **geophysical and geological models** for understanding geoneutrino signal.



# Mantle geoneutrinos (KL)

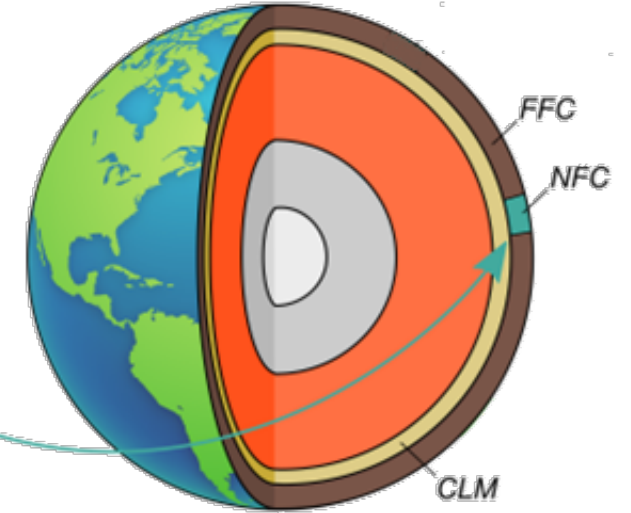
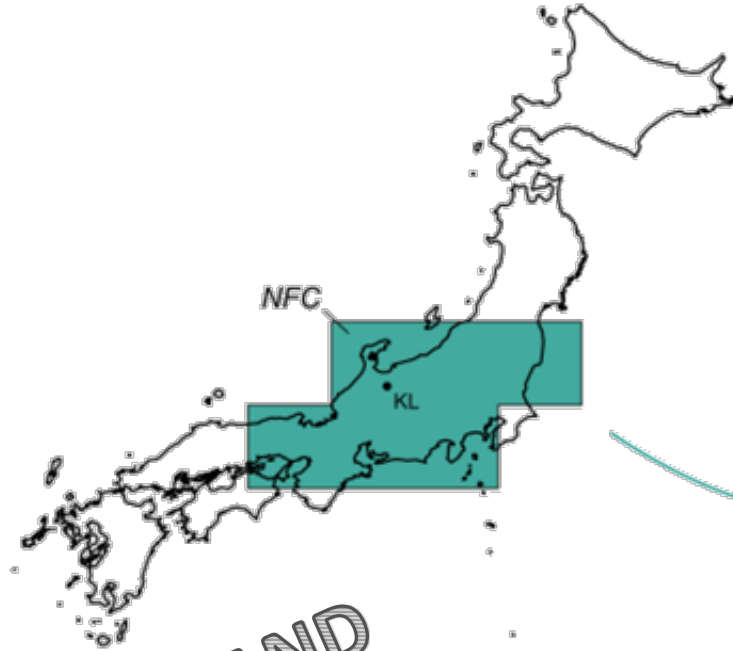
$$S_M^{KL}(U + Th) = S_{Exp}^{KL}(U + Th) - S_{NFC}^{KL}(U + Th) - S_{FFC}^{KL}(U + Th) - S_{CLM}^{KL}(U + Th)$$

	KamLAND
$S_{NFC}(U+Th)$ [TNU]	$17.7 \pm 1.4$
$S_{FFC}(U+Th)$ [TNU]	$7.3^{+1.5}_{-1.2}$
$S_{CLM}(U+Th)$ [TNU]	$1.6^{+2.2}_{-1.0}$

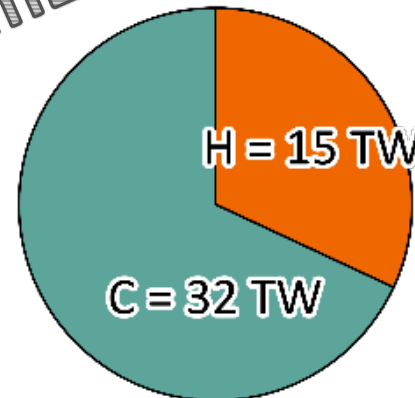


$$S_M^{KL}(U+Th) = 4.8^{+5.6}_{-5.9} \text{ TNU}$$

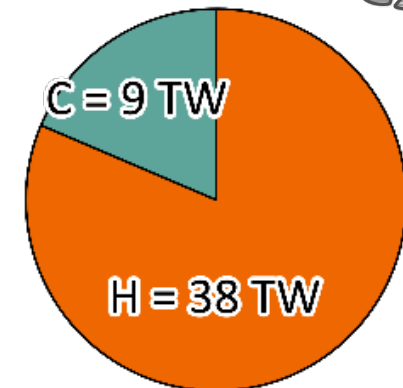
$$S_M^{BX}(U+Th) = 21.2^{+9.6}_{-9.0} \text{ TNU}$$



KamLAND



Borexino

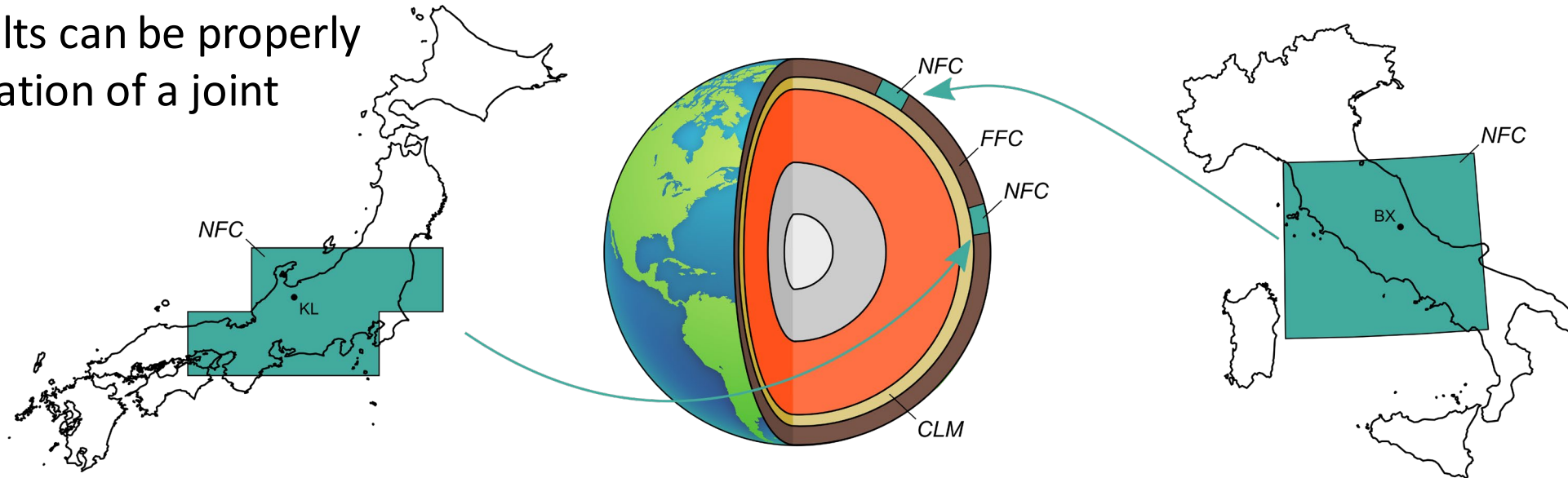


# Combined mantle geoneutrinos (KL + BX)

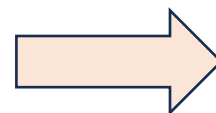
$$S_M^{KL}(U + Th) = S_{Exp}^{KL}(U + Th) - S_{NFC}^{KL}(U + Th) - S_{FFC}^{KL}(U + Th) - S_{CLM}^{KL}(U + Th)$$

$$S_M^{BX}(U + Th) = S_{Exp}^{BX}(U + Th) - S_{NFC}^{BX}(U + Th) - S_{FFC}^{BX}(U + Th) - S_{CLM}^{BX}(U + Th)$$

- Assuming the site-independence of the mantle signal, the results can be properly combined in the estimation of a joint bivariate PDF.



	KamLAND	Borexino
$S_{NFC}(U+Th)$ [TNU]	$17.7 \pm 1.4$	$9.2 \pm 1.2$
$S_{FFC}(U+Th)$ [TNU]	$7.3^{+1.5}_{-1.2}$	$13.7^{+2.8}_{-2.3}$
$S_{CLM}(U+Th)$ [TNU]	$1.6^{+2.2}_{-1.0}$	$2.2^{+3.1}_{-1.3}$



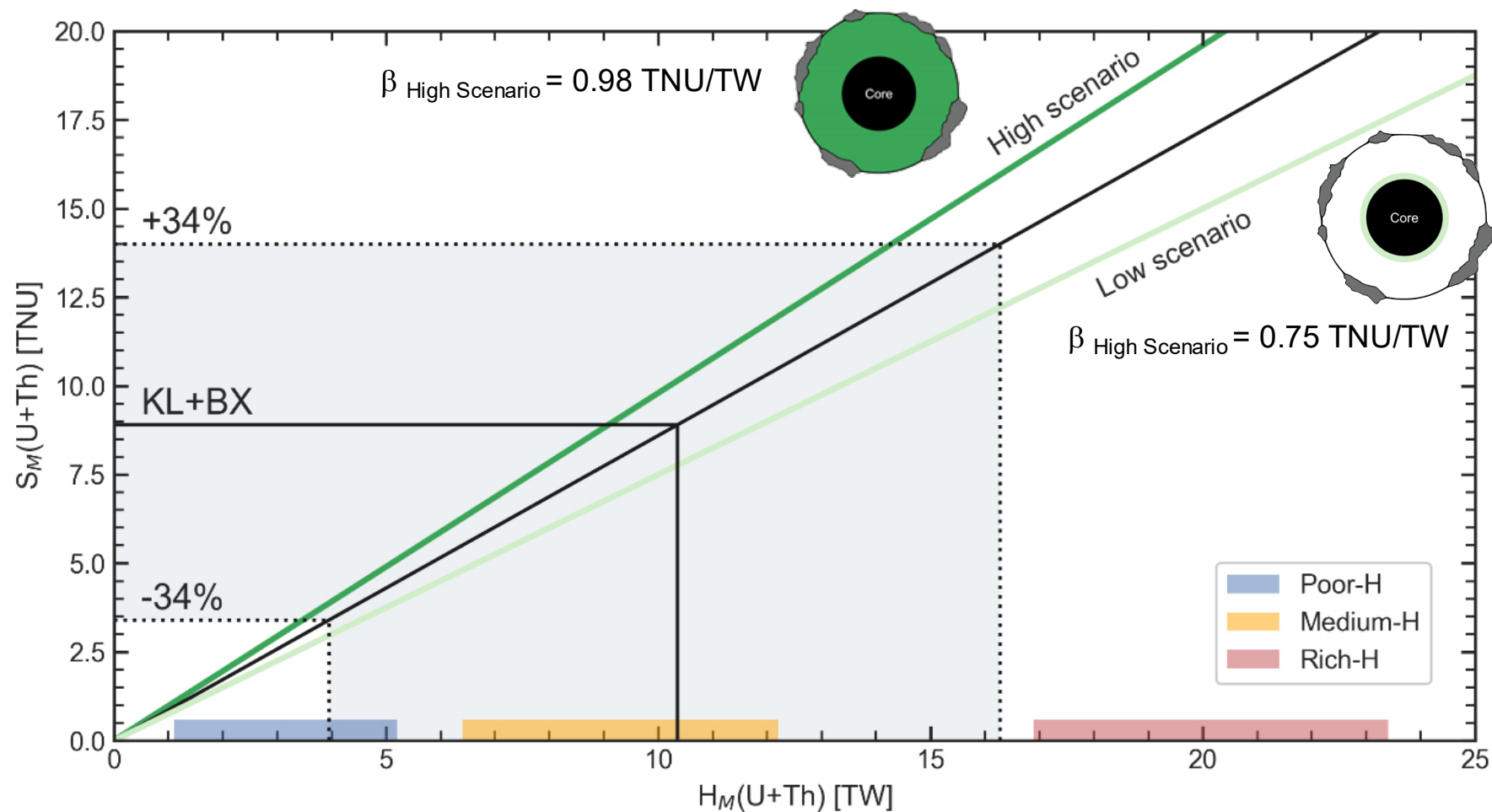
$$S_M^{KL+BX}(U+Th) = 8.9^{+5.1}_{-5.5} \text{ TNU}$$

For more details see Fabio's talk



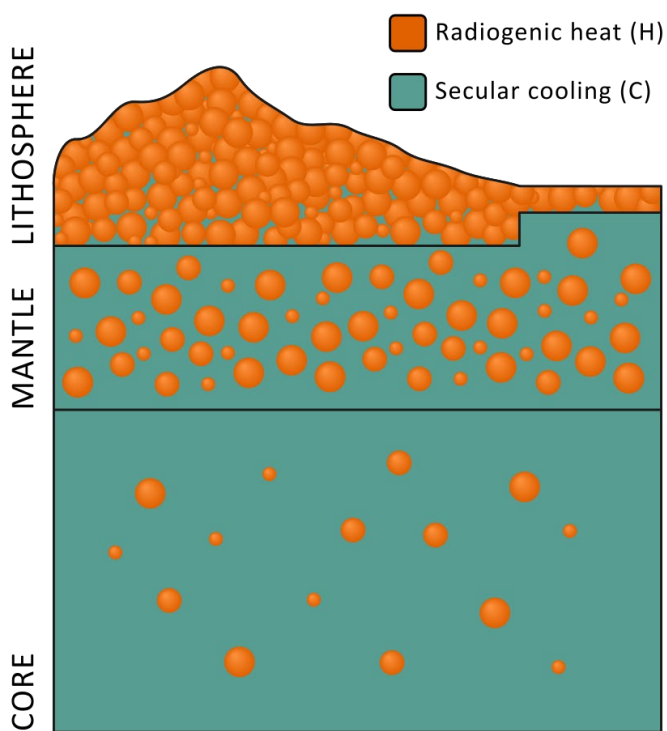
# Mantle radiogenic power (KL+BX)

Class	References
Poor – H	Jackson and Jellinek, 2013
	O'Neill and Palme, 2008
	Javoy and Kaminski, 2014
	Javoy et al., 2010
Medium - H	McDonough and Sun, 1995
	Lyubetskaya and Korenaga, 2007
	Palme and O'Neill, 2007
	Arevalo, 2010
	Wang et al., 2018
	Palme and O'Neill, 2014
Rich - H	Turcotte, 2002
	Turcotte, 2014



	Poor	Medium	Rich	KL+BX
$H_M(\text{U+Th}) [\text{TW}]$	$3.2^{+2.0}_{-2.1}$	$9.3 \pm 2.9$	$20.2^{+3.2}_{-3.3}$	$10.3^{+5.9}_{-6.4}$

# Earth's heat (BX + KL)



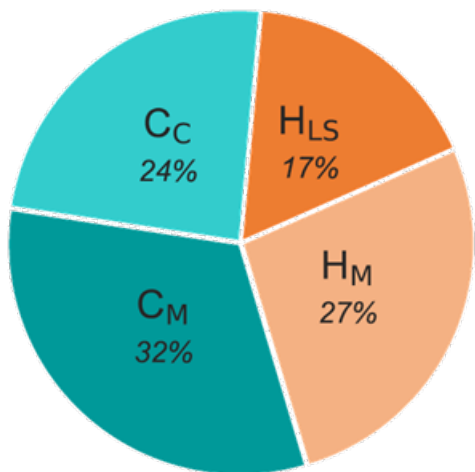
$$C = Q - H$$

$$C_M = Q - H - C_C$$

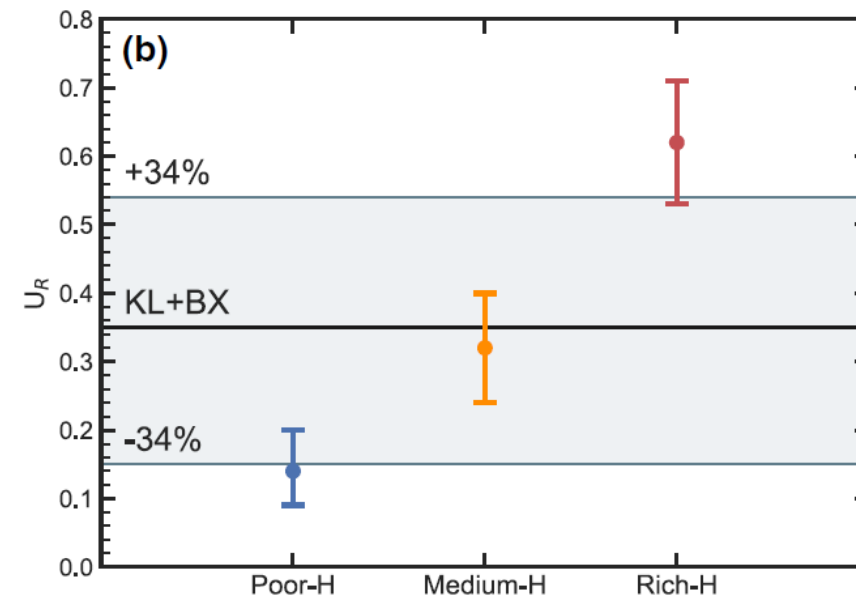
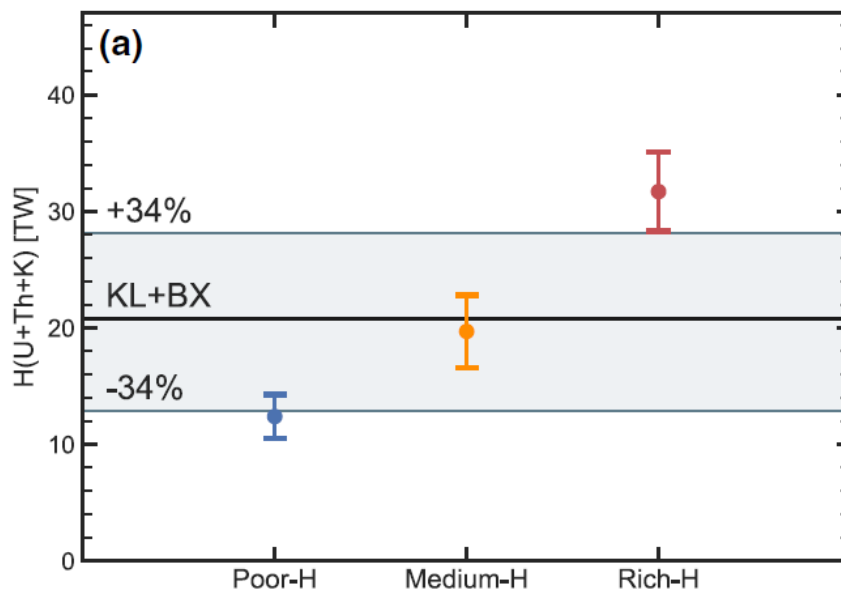
$$H_M = H - H_{LS} - H_C$$

$$H_{LS} = H_{CC} + H_{OC} + H_{CLM}$$

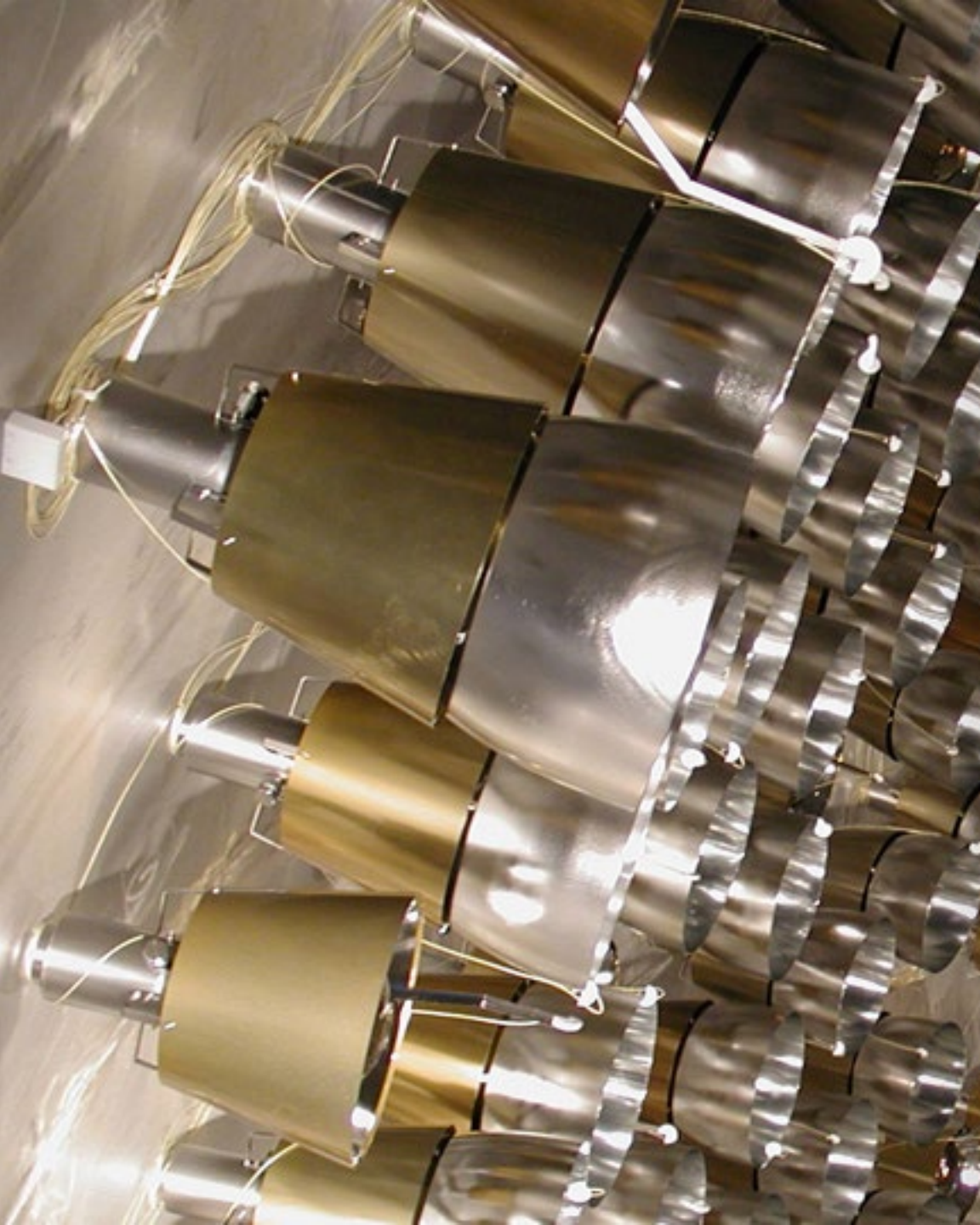
$$U_R = \frac{H - H_{CC}}{Q - H_{CC}}$$



	Combined KL + BX
<b>Q [TW]</b>	<b>47 ± 2</b>
$H_{LS}(U+Th+K)$ [TW]	8.1 <sup>+1.9</sup> <sub>-1.4</sub>
$H_M(U+Th+K)$ [TW]	12.5 <sup>+7.1</sup> <sub>-7.7</sub>
<b>H [TW]</b>	<b>20.8<sup>+7.3</sup><sub>-7.9</sub></b>
<b>C [TW]</b>	<b>26 ± 8</b>





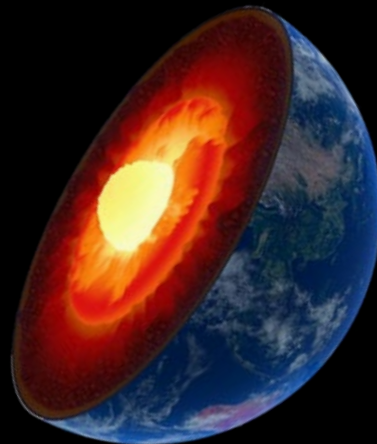


# Take-away messages

- To deeply understand the experimental geoneutrino results, the use of refined geological models is essential
- The Borexino (KamLAND) observations favor geological models that predict a relatively high (low) concentration of radioactive elements in the mantle
- The combined mantle measurements (BX + KL) falls within the prediction of the Medium-H models
- The era of "multi-site detection" of geoneutrinos is definitely open...

## *References:*

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Thank you

